ASSESSMENT OF RIVER NZOIA BASIN MORPHOMETRIC CHARACTERISTICS AND FLOODING RISKS IN BUDALANG'I AREA, BUSIA COUNTY

BY

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DECLARATION

1, Celestine Agnes Nabwire declare that this research project is my original work and has not been presented for the award of a degree or diploma in any other university, it contains no material previously published or written by any other person except where due reference is made in the work itself.

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DEDICATION

To my parents: Mrs. Regina Nekesa Ouma, my Late Dad Mr. Gabriel Ouma Juma and my cousin Mr. Norbert Abachi

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I would like to express my gratitude towards my Mum and my relatives for their support.

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ABSTRACT

Flooding continues to be a common environmental hazard in both developed and developing countries. River Nzoia experiences perennial flooding in its lower reaches especially in Budalangi area of Busia county. The mean annual discharge of the River Nzoia is estimated at 1777Mm3 /year. From a physiographic and land use point of view the basin has four distinct zones: a mountain zone, plateau zone, transition zone and lowland zone. The mountain zone is forested but suffers severe land degradation; the plateau zone is the major farming zone. Small scale farming continues in the transition and flood prone lowland areas. The flood prone area is generally flat and swampy. There are two rainfall peaks in the catchment; the first peak comes in the months of April to June, while the other occurs in July to September. Comparatively to other parts in Kenya, the basin receives high rainfalls, whose average annual values vary between 1,000 to 1,500 mm. Several studies (JICA, 1987; JICA, 1992; APFM, 2004; ADCL, 2006) have been carried out in Nzoia basin with a view to characterizing the basin on the basis of climatic data and land use information but none has incorporated landform parameters because geomorphologic studies have not been attempted. River morphometric information is very important in any undertaking to control incidence of flooding in an area. In Kenya Few studies have been carried out on River morphometrics in relation to flooding. This study therefore seeks to analyze the characteristics of River Nzoia basin and how they contribute to flooding of Budalang'i flood plains. The study was guided by the following specific objectives; to examine the drainage basin morphometrics characteristics of the River Nzoia, to determine the influence of rainfall intensity and duration on flooding in Budalang'i flood plains and to assess approaches of managing floods in Budalang'i flood plains. Descriptive research design was used for the study. The target population in this study was drawn from two components of the unit of analysis, namely the geophysical (River Nzoia basin) and the social (households). Simple random sampling technique was used to sample the respondents for the study. A sample size of 71 respondents was sampled from the locations along River Nzoia in Budalang'i flood plains. Interviews and observation were used as methods of data collection. Qualitative data analysis techniques were used to analyze the data where thematic and content analysis was carried out based on the study objectives. The major datasets that were used for this study include Landsat Satellite Imageries, Digital Elevation Data (DEM) and Auxiliary data that included Roads, administrative boundaries, ground truth data for land use land cover classes and topographic map of the area. The study found out that high volumes in upper catchment areas during heavy rains of long rain seasons are at higher velocity due to sudden change of river gradient. This result in higher volumes received at the lower catchment areas known as flood plains (in-flow) against that volume expected to be released to the lake (out flow). This volume difference causes the river to break its normal banks in the lower zones. Land use land cover analysis between 1986-2018 shows that there is an increase in size of the bare ground area which is an indication of wetland degradation and a consequent replacement of wetland vegetation with bare soil that is prone to erosion. Increased farming activities along the Nzoia River has resulted to deforestation in search of cultivation areas increasing the vulnerability of the area to flooding. The rapid expansion of urban centres in the lower parts of the catchment (Rwambwa, Kakamega, and Bungoma) can be said to be a major contributing factor to the annual devastating floods. Based on River morphometry, the study concluded that there is no consistent relationship between relief ratio and basin order in this basin due to high variance in relief on the divide line. However, high relief ratios of above 0.03 are observed as characteristics of upland sub-basins and low value of about 0.01 for sub-basins at the lowlands. This change has explained the flooding phenomena at the lowlands. The computation of bifurcation ratio shows that the values ranged

from 1.5 and 3.4 which are within Strahler's range. The average value of Bifurcation ratio is 2.2 which is closer to the lower value of 2. Based on bifurcation ratio, River Nzoia basin flows may encounter delayed time to peak and this is a good characteristic for planning and conveying flood forecasts. The drainage pattern of this basin can be classified as fern-shape. This presents a lower concentration time, and it generates higher flow. The major flooding risks that were identified during the study include; Displacement of households, Damage to Farmlands, Damage to Bridges and Roads, Disruption of learning programmes, Economic and Environmental aspects. Approaches to flood management will include both Community non-structural measures and Structural measures for flood control. The study recommends that there is need to undertake further research in land-use land cover changes, river morphometrics and review of the existing flood control and /or mitigation measures

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ABBREVIATIONS AND ACRONYMS

DEFRA	Department for Environment, Food and Rural Affairs
EC	European Commission
ENSO	El Nino Southern Oscillation
FRC	Flood Resilient City
GHA	Greater Horn of Africa
LBDA	Lake Basin Development Authority
SSA	Sub-Saharan Africa
SWMP	Surface Water Management Plans
LULC	Land Use Land cover changes
RCMRD	Regional center for mapping and resource development

CHAPTER ONE: INTRODUCTION

1.0 Introduction

This chapter gives the background, objectives and the concepts studied.

1.1 Background of the study

One of the global most environmental natural occurrences is river flooding. Flooding is a temporal condition of partial or complete inundation of normally dry areas from overflow of inland or tidal waters or from unusual and rapid accumulation or runoff" (Jeb & Aggarwal, 2008). As such, river floods occur when water flows over the river banks into the adjacent land also known as flood plains. Floods are triggered by anthropogenic activities such as increase in settlement areas, expansion of agricultural areas, deforestations, alterations of natural drainage and river basin patterns, and climate change (EC, 2007). Naturally, flooding can involve overflow of a river because of prolonged seasonal rainfall, rainstorm, snowmelt, dam-breaks, accumulation of rainwater in low-lying areas with a high-water table, or inadequate storm drainage. In addition, floods can be caused by intrusion of sea water onto coast lands during cyclonic/tidal surges (Lidstone & DeChano, 2004).

From a purely geomorphic perspective, high flows become floods when they provide large quantities of sediments or alter the morphology of the river channel and floodplain. From a human perspective, high flows become floods when they cause injury or death of people, or when they result in relatively minor damage on property, belongings or source of livelihood. Small floods result in relatively minor damage, but they may have higher cumulative cost because they are frequent and occurs in many locations. Larger floods are rare and have the potential to cause heavy loss of life and economic damage (WMO, 2006.)

According to Emergency Events Data Base (EM-DAT), floods comprised two-fifths of all disasters in 2015 and were responsible for 25 per cent of the total economic damage and 37 per cent of the total disaster affected population in Asia Pacific. EM-DAT further asserts that floods occurred 63 times resulting to 1863 human deaths and affecting about 21.7 million people. EM-DAT estimated an economic loss of about US\$11.5 billion to the Asia Pacific countries. In South America, Brazil alone has experienced a number of major floods in the past 30 years, in January 2011 floods killed 900 people and lead to a significant economic loss of approximately 1 billion US dollar (Whitney *et al.* 2013).

In Fiji a great flood in 1993 brought on damages adding up to \$ 100 million, 23 lives were lost, and more than 120,000 individuals suffered losses thereby, ruinously impacting the national GDP and the government's advancement strategies and programs. Assets reserved for capital improvement works were immediately diverted for help and recovery.

In Africa, it is estimated that floods fatalities and the accompanying economic losses have increased over the past half-century. In 2007 floods displaced more than a million people in Burkina Faso, Togo, Mali and Niger, Uganda, Ethiopia, Sudan and claimed over 500 lives; while in 2009 torrential rains resulted to floods that affected 600,000 people in 16 West African nations, with the worse damages being recorded in Burkina Faso, Senegal, Ghana and Niger (United Nations, 2009).

Conway (2002) conducted a study in sub-Saharan Africa and he noted that rainfall and river flows in Africa displays high levels of variability across a range of spatial and temporal scales. Thought out Africa, this variability brings significant implications for society and causes widespread acute human suffering and economic damage. Examples of such variabilities include prolonged periods of high flows for rivers draining large parts of East Africa and multi-decadal anomalies in river flows regimes in parts of west Africa where long-term mean yields of freshwater into Atlantic Ocean fell by 18% between 1951-70 and 1971-89 (Mahe' and Olivry).

Floods are the leading hydro-meteorological disaster in East Africa. In Kenya, floods are emerging as the most prevalent climatic disaster (RoK, 2007; ISDR). Perennial floods affect low-lying regions of the country such as river valleys, swampy areas, lakeshores and the coastal strip that are unevenly distributed in the country's five drainage basins. Geographically, the western, northern, eastern, central and south-eastern parts of the country are quite susceptible to seasonal floods in the wet seasons of March-April May (MAM) and October-November-December (OND).

The Lake Victoria Basin in western Kenya is the most flood-prone region in the country (RoK, 2007). Western Kenya is characteristically wet throughout the year with no distinctive dry season. High rainfall is received in the months of March to September, with significantly lower rainfall in January and February (SoK, 2003 in WRI et al., 2007). Contrary to the bimodal rainfall pattern in the rest of the country, a third rainfall season is experienced during the cold and dry months of

June-July-August (JJA). (ICPAC, 2007). The mean annual rainfall in western Kenya is above 1600 mm (WRI et al., 2007).

The prevalence rates of floods in Kenya stands at 27% and affects 5% of the population affected by disasters. Floods related fatalities constitute a whopping 60% of disaster victims in Kenya (UNEP, 2009). Flood occurrence trends in western Kenya is increasingly becoming a major concern to the country's socio-economic development due to the substantial economic and financial losses incurred to respond to frequent flood disasters.

Riverine floods are the most dominant floods in Kenya and mostly occurs along floodplains as a result of exceeded stream flow capacity leading to over spilling pf the natural banks or artificial embankment (smith and Ward,1998). River that most experience flooding in kenya include: Nzoia, Nyando, Tana, Athi, Yala, Sondu, Nairobi and Kuja. River floods in western kenya have high frequency as compared to other parts of the country, making the communities of the region vulnerable to flood risks especially frequent floods of River Nzoia in Budalangi floodplains

1.2 Statement of the problem

Natural disasters cause much misery, especially in developing countries where they cause great stress among low-income economies. Approximately 70 per cent of all global disasters are linked to hydrometeorological events. Flooding poses one of the greatest natural risks to sustainable development. Flood losses reduce the asset base of households, communities and societies through the destruction of standing crops, dwellings, infrastructure, machinery and buildings, quite apart from the tragic loss of life. In some cases, the effect of extreme flooding is dramatic, not only at the individual household level, but in the country. The global economic losses from flooding exceeded \$19 billion in 2012 and are rising rapidly globally (Munich Re 2013).

Flooding continues to be a common environmental hazard both in developed and developing countries, like Kenya. Many parts of Kenya experience unexpectedly long rainfall in mid-April to May and short rains in September to November. During the long rains, the Lake Victoria Basin, which comprises of the Nzoia River basin among the other river basins busts its banks resulting to floods.

The western region of Kenya is one of the regions most influenced by floods. Frequent flooding occurs in parts of western Kenya particularly Budalang'i and Bunyala plains in the Nzoia river

basin. The Nzoia basin has a high incidence of poverty with a rich natural resource endowment. While the Western Kenyan region is endowed with natural resources such as forests, rivers and lakes, which should be adequate for poverty reduction, and vulnerability nonetheless afflict many in the region. The communities in the Nzoia basin are confronted with flooding, disease, and degradation of natural resources, especially land. The situation is aggravated by perennial flooding, mismanagement of natural resources, and the HIV/AIDS pandemic.

Year	No of people affected
1997-1998	widespread
2004	12,000
2006	15,888
2007	27,000
2008	10,000
2010	200
2011	58,000
2013	12,000
2016	7,000

Table 1: Flooding history in Budalangi area

Source: KNBS 2018

The physical setting of Budalangi at the floodplain of Nzoia River and increased runoff from degraded catchments are contributory factors to the flooding. This implies that the policy measures that have been instituted by the government to mitigate the problem have had dismal impact in the Budalangi area. The degradation of the catchment is reflected in its sediment loading and deposition into Lake Victoria which has seen the morphology of the coastline at the mouth of Nzoia River and the aerial coverage by water in the lake change over the years. Attempts to control floods in this region by construction of dykes have not successfully solved the problem.

Assessment of flood risk is basis for decision-making in flood management at international, national, regional and local levels. To curtail the impacts of floods on people's social and economic aspects, it is imperative to understand the role of the river morphometric characteristics on flooding. Langbein *et al.* (1944) asserts: "the structure of drainage basins is a sensibly permanent characteristic which influences mainly the concentration or time distribution of the discharge from a drainage basin

Despite the concerted efforts by the Ministry of Water and Irrigation (MWI) to tame the floods by the rehabilitation and raising of the dyke embankments, the solution has been elusive. (Mutiso 2007)

Previous studies in Budalangi focused on; Flood Disaster Preparedness, household and community experiences and perceptions on climate change impacts due to floods but the morphometric aspects of the river Nzoia basin in relation to flooding in Budalangi has been scantly studied.

In Kenya, River morphometry studies are few and thus scanty. Nzoia basin is no exception. Therefore, this study will analyze the morphometric characteristics of River Nzoia basin and how it contributes to flooding of Budalang'i flood plains as part of the early warning for floods and additional knowledge on the ever-changing flooding characteristics.

1.3 Research questions

- i. What is the effect of rainfall intensity and duration on flooding in Budalang'i flood basin?
- ii. How does River Nzoia basin morphometry affect the flooding in Budalang'i floodplains?
- iii. What are the suitable approaches for managing floods in Budalang'i?

1.4 Research objectives

The overall objective of the study is to assess River Nzoia basin morphometry and flooding risks in Budalang'i flood plains.

The specific objectives are:

- i. To determine the effect of rainfall intensity and duration on flooding in Budalang'i flood basin
- ii. To assess the effect of basin morphometry on flooding of Budalang'i flood plains
- iii. To assess approaches of managing floods in Budalang'i flood plains.

1.5 Justification of the study

Flooding has been a persistent occurrence in Bunyala Sub County. The flooding impact in the area has been immense despite modern and traditional mitigation approaches that have been taken in addressing the flooding menace in the area. The government's efforts to form task forces for the mapping of flood prone areas and institutionalize Disaster Management Committees are yet to produce any tangible results. This might be associated with modern technology on flood

management that has overlooked opinion and experience of the local residence as well as paying minimal or no attention to the historical information likely to suggest probable factors that trigger flooding

Nzoia River basin is the largest sub-basin in the Lake Victoria North basin of Kenya and yet, the geomorphology as a component of the basin's general physical characteristics and its linkage to flooding has scantly been studied. River morphometric characteristics are dynamic and require regular documentation in order to understand the flooding risks.

Recent government proposal to manage flood hazard in the basin by adoption of integrated flood management strategy (ADCL, 2006) will need a comprehensive characterization of the basin on the basis of all variables. The communities who live in the lowland zone of River Nzoia continue to experience annual flood disasters due to failure of structural flood mitigation systems like earth dykes in controlling high flows. Destruction of natural forest cover in the middle and upper zones of Nzoia catchment continue unabated. The linkage between geomorphology and increased incidences of flood disaster is poorly understood amongst planners and communities occupying the basin. Therefore, the objective of this study is to establish geomorphic characteristic of Nzoia basin with reference to a few commonly used morphometric parameters of relief ratio, bifurcation ratio, and drainage density. The researcher understand that the full scope of geomorphology goes beyond these three parameters but assumes that the four parameters will provide a basic geomorphic understanding of the area for planners.

1.6 Scope and the limitation of the study

The study focused on assessing morphometric characteristics (stream order, stream number, stream length, Burfacation ratio, Relief ratio, and Drainage density) within the catchment, rainfall intensity and duration and land use Landcover changes as factors responsible for flooding and how they affect the individual households. The study also identified and quantified key causes and accelerators of flooding in River Nzoia catchment and assess seasonal variation of the same sinks in terms of material composition, size and permanence as well as the interrelationship between these sediment sinks and vegetation. Due to time and financial constraints field assessment and measurements was not done to cover both the dry and rainy seasons.

1.7 Definition of Terms and concepts

Floods-an overflow of a large amount of water beyond its normal limits, especially over what is normally dry land.

Flooding-a situation in which an area is covered with water, especially from rain

Basin Morphometry - is the mathematical measure of the shape and dimensions of the landform's geomorphic conditions and hydrological processes

Disaster-A disaster is a serious disruption, occurring over a relatively short time, of the functioning of a community or a society involving widespread human, material, economic or environmental loss and impacts, which exceeds the ability of the affected community or society to cope using its own resources

Early Warning-An early warning system warns people that something bad is likely to happen

Flood prone areas-The regions along a waterway inclined to flooding and can be separated into various zones, specifically the floodplain and floodway. (ISDR 2004; Wright 2008).

Watershed- an area or ridge of land that separates waters flowing to different rivers, basins, or seas.

Risk – hazard is the likelihood that destructive outcomes or misfortune can happen because of connections amongst normal and anthropogenically actuated dangers (ISDR 2010a). Hazard can likewise be characterized as an element of risk, powerlessness and adapting limit (Baas et al. 2008; Botha & Louw 2004; Boudreau 2009; Hossini 2008)

Mitigation- The reducing or constraint of the antagonistic effects of risks and related fiascos (UNISDR, 2007)

CHAPTER TWO: LITERATURE REVIEW

2.0 Introduction

In this section, Basin morphometric, Rainfall Intensity and duration and flood management and flood control aspects are discussed.

2.1 Basin morphometric and flooding

Morphometry refers to the geographical articulation of land by method for zone, incline, shape, length, and so forth (Horton, 1932; Clarke, 1966). Therefore, the form and structure of drainage basins and their associated drainage networks can be described by their morphometric parameters. These parameters influence catchment stream flow design through their effect on fixation time (Jones, 1999; Ajibade, Ifabiyi, Iroye, and Ogunteru, 2010). These landscape parameters are huge on the grounds that stream flow can be communicated as a general capacity of geomorphology of a watershed. Basins on Flood Vulnerability in Makurdi Town, Nigeria still stand legitimate after (Jones, 1999; Jain and Sinha, 2003; Okoko and Olujinmi, 2003), announced that the morphometric investigation of water basin gives a rich portrayal of the scene, as well as fills in as a ground-breaking method for looking at the frame and procedure of seepage basin that might be generally isolated in space and time.

Morphological studies of rivers are very important to study the behavior of a river, its aggradations/degradation, shifting of the river course, flooding and erosion of river bank and to plan remedial measure for erosion, flooding and other related problems (Oruonye, Ezekiel, Atiku, Baba & Musa, 2016). A research done in Nigeria's Kereke and Ukoghor basins showed the influence of morphometric variables on flooding in drainage basins using GIS. Okoghor drainage basin morphometric variables caused high peak flow and shorter duration concentration time thereby inducing high magnitude flood within the basin compared with morphometric properties of Kereke drainage basin (Oyatayo, Bello, Ndabula Godwill & Ademola, 2017). The influence of morphometric influence is however encouraged by other basin characteristics such as climate, soil features and land use pattern which acts as contributory factors to flooding in drainage basins (Oyatayo, Bello, Ndabula Godwill & Ademola, 2017).

The soils of the floodplain in the lower reaches of the Nzoia River are all alluvial. The river meanders in the floodplain depositing silt during seasonal floods. The major parts of the floodplain contain black cotton soils, while other areas have coarse textured sand silt mixture. In some places, there exist saline soils (Oyugi et al. 2003).

The upper Nzoia Basin contains extensive seasonal swamp areas in the highland, medium rainfall zones that are mainly utilized for grazing due to poor drainage. Once the river bursts its banks, the resultant flooding displaces close to 30,000 persons every rainfall season. People are usually evacuated to safer higher grounds during floods. Entire homesteads are swept away, property and crops worth hundreds of thousands of shillings are lost and many people perish as rivers break their banks, rendering large areas of land inaccessible. The floods enhance poverty, because crops and businesses are destroyed, (Oyugi et al. 2003).

The lower reaches of River Nzoia are flat lands with low gradients. All the six locations in Budalangi Division are perennially affected by flooding. These are; Bunyala East, West, South, North and Central and Khajula location. Traditionally, the people have settled in the flood plain and this has meant that people are directly affected by the perennial flooding. The April 2003 floods displaced 4000 families or about 24,000 people, covering an area of 60 km² with people moved to camps on higher areas and relying on food relief (Paolo et al. 2013)

This situation has arisen after breaching of protection dykes that were constructed by the Ministry of Water from 1977 to 1984. The works involved construction of 32.8 km of earthen dykes (16.6 km in the southern side and16.2 km) on the northern side that required a replacement of 690,000 m3 of soil. They were designed for 25-year flood protection of 750 m³/s. Now the dykes are occasionally overgrown with vegetation, breached at 20 points and have outlived their life of 20 years. The budgetary maintenance requirements are therefore high and have been diminishing annually.

A report given to the government in 1992 after the assessment of the area proposed an irrigation scheme as the solution to the unending flood menace. In the report, it was proposed that a dam be constructed to trap the excess water that overflows the river. This would be followed by the construction of two canals, one in the northern riverbank measuring 40 km and the other on the southern side measuring 35 km, draining water onto the flat farms for irrigation. 'The project

could cost about \$20 million if implemented'. It is evident that the present course of the river is highly unstable and could change from one moment to the next after a flood, if no protective works are provided (GoK, 2002)

The 2003 floods in Budalang'i saw nearly 24,000 out of a population of 56,000 people displaced (IFRC, 2003). Some 10,000 people were accommodated in the District Officer camp, necessitating health emergency measures to control possible outbreaks of waterborne diseases. Scarcity of water sources and the contamination of pipes and borewells aggravated an already acute problem. Fishing and farming are the major economic activities in Budalang'i. Due to the floods and concentration in the camps, the economic activities are greatly disrupted

2.1.1 Drainage density (Dd)

Drainage density has for quite some time been perceived as topographic normal for crucial noteworthiness. This emerge from that reality that drainage density is touchy parameter which from numerous points of view gives the connection between the shape characteristics of the basin and the procedures working along stream course (Gregory and Welling, 1973). It mirrors the land use and influences penetration and the basin reaction time between precipitation and release. It is additionally of geomorphological intrigue especially for the advancement of inclines. Drainage basin with high Dd shows that an extensive extent of the precipitation keeps running off. Then again, a low drainage density demonstrates the most precipitation penetrates the ground and few channels are required to convey the spillover

(Roger, 1971). Dd is viewed as a vital file; it is communicated as the proportion of the aggregate total of all channel fragments inside a basin to the basin territory i.e., the length of streams per unit of drainage density. It is a measurement reverse of length (Horton, 1932). Dd is a proportion of the surface of the system and demonstrates the harmony between the erosive intensity of overland flow and the opposition of surface soils and shakes. The factors influencing drainage density incorporate geography and density of vegetation. The vegetation density impacted drainage density by restricting the surface layer and moderates down the rate of overland flow and stores a portion of the water for brief timeframes. The impact of lithology on drainage density is stamped. Porous rocks with a high penetration rate decrease overland flow, and therefore drainage density is low.

The drainage density exerts on flood peaks significant controls which can be broadly divided between direct and indirect effects (Merz and Bloschl, 2008; Bloschl, 2008). Among the most significant direct effects there is the control associated with the length of the stream network and hillslope paths. Because flow velocity is higher in the river network, Dd significantly affects the concentration time and therefore the peak flow magnitude. It follows that an increasing drainage density implies increasing flood peaks. Moreover, a long concentration time implies more opportunities for water to infiltrate. Therefore, a decreasing Dd generally implies decreasing flood volumes.

Murphey et al. (1977) analyzed the relationship among flood hydrograph characteristics, like rise time, base time, mean peak discharge and flood volume, and basin parameters such as, among the others, contributing area and drainage density. Plaut Berger and Entekhabi (2001) and Humbert (1990) proved that drainage density is significantly related to the basin runoff coefficient. Gresillon (1997) studied 100 African catchments and concluded that Dd and watershed area are two very important variables to explain the shape of the recessing limb of the hydrograph, while in tropical regions Dd is effective on the runoff coefficient during flood events. Yildiz (2004) carried out a numerical simulation showing that streamflows simulated by a rainfall-runoff model steadily increase with increasing Dd. Recently, Grauso et al. (2008) showed that Dd is a fundamental explanatory variable in regression models for estimating the annual amount of suspended sediment yield. The explaining capability of Dd was found to be extremely relevant, therefore suggesting that a significant link indeed exists between Dd and the river flow regime. A recent and interesting review is provided by Wharton (1994) who studied the usefulness of Dd in rainfall-runoff modelling and runoff prediction.

2.1.2 Stream order (U)

There are four different system of ordering streams that are available Gravelius (1914), Horton (1945), Strahler (1952) and Schideggar (1970). Strahler's system, which is a slightly modified of Hortons system, has been followed because of its straightforwardness. Where the smallest, unbranched fingertip streams are designated as 1st order, the confluence of two 1st order channels give a channel segments of 2nd order, two 2nd order streams join to form a segment of 3rd order and so on. When two channels of different order join then the higher order is maintained. The trunk stream is the stream segment of highest order.

2.1.3 Basin shape

The shape of the basin mainly directs the rate at which the water is provided to the main channel. Shape will contribute to the speed with which the runoff reaches a river. A long thin catchment will take longer to drain than a circular catchment. A circular shaped drainage basin leads to rapid drainage whereas a long drainage basin will take time for the water to reach the river. The main indices used to analyze basin shape and relief is the relief ratio.

2.1.3.2 Relief ratio

In accordance to Schumm (1956), the maximum relief to the horizontal distance along the largest dimension of the basin parallel to the principal drainage line is termed as relief ratio. It computes overall steepness of the drainage basin. Relief ratio is an indication of the intensity of erosional process operating on slope of the basin (Schumm, 1956). On the basis of work carried out by Shumm (1963), it is dimensionless height-length ratio equal to the tangent of angle formed by two planes cross at the mouth of the basin, one exhibit the horizontal and the other passes through the highest point of the basin (Singh, 1997). When the basement rock of the basin is resistance and low degree of slope heralds or give rise to low value of relief ratio (Gottschalk, 1964) Relief ratio is inversely proportional to the drainage area and the size of given drainage.

2.1.4 Stream length (SI))

According to Horton (1945), the aggregate average lengths of stream segments of each of the consecutive orders in a catchment tend closely to estimate a direct geometric series in which the first term is the mean length of streams of the first order. The total length of singular stream sections of each order is the stream length of that order. Stream length measures the average (or mean) length of a stream in each order and is computed by dividing the total length of all streams in a particular order by the quantity of streams in that order. The stream length in each order rises exponentially with increasing stream order.

2.1.5 Bifurcation ration (Rb)

Bifurcation Ratio is the number of streams in a low order to the number of streams in the next high order (Horton, 1945). The bifurcation ratio varies from a minimum of 2 in "gently sloping drainage basins" to 3 or 4 in "mountainous or highly dissected drainage basins"; it is a parameter used in equations giving the number of streams in a basin. The higher the bifurcation ratio, the shorter the

time taken for discharge to reach the outlet, and higher was the peak discharge – leading to a greater probability of flooding. Bifurcation ratio correlates positively with drainage density i.e., a high bifurcation ratio indicates a high drainage density. Higher Bifurcation ratios also suggest that the area is tectonically active. The bifurcation ratio can also show which parts of a drainage basin is more likely to flood, comparatively, by looking at the separate ratios. The bifurcation ratio varies from a minimum of 2 in "flat or rolling drainage basins" to 3 or 4 in "hilly or highly dissected drainage basins"

2.1.6 Stream Number (Nu)

The total number of stream sections existent in each order is the stream number (Nu). Nu is number of streams of order u. The total number of stream sections is found to decline as the stream order rises in most river basins Horton (1945).

2.1.7 Slope

Slope is very important in how quickly a drainage channel will convey water, and therefore, it influences the sensitivity of a watershed to precipitation events of various time durations. Watersheds with steep slopes will rapidly convey incoming rainfall, and if the rainfall is convective (characterized by high intensity and relatively short duration), the watershed will respond very quickly with the peak flow occurring shortly after the onset of precipitation. Steep slopes tend to result in rapid runoff responses to local rainfall excess and consequently higher peak discharges (WWAP, 2006). The runoff volume is also affected by slope. If the slope is very flat, the rainfall will not be removed as rapidly. The process of infiltration will have more time to affect the rainfall excess, thereby increasing the abstractions and resulting in a reduction of the total volume of rainfall that appears directly as runoff (WWAP, 2006).

2.2 Rivers flooding

Asia continent is much affected by floods and countries like India, China, Philippines, Iran, Bangladesh and Nepal are extremely vulnerable (WWAP, 2006). It indicates that the majority of flood disasters' victims are poor people of developing countries, who suffer most and are the first casualty of such incidents. High relief, steep slopes, complex geological structures with active tectonic processes and continued seismic activities, and a climate characterized by great seasonality in rainfall, all combine to make Nepal natural disasters prone area, mainly earthquake, water induced disaster such as flood, landslides, glacier lake outburst flood among others. About

6000 rivers and rivulets drain Nepal, with a combined total length exceeds 45,000 km, having a total drainage area of 191,000 sq. km. They discharge more than 200 billion m³ of water annually. As the river emerge into plain from steep and narrow mountains gorges, they spread out with an abrupt gradient decrease that has three major consequences; deposition of bed load, changes in river course, and frequent floods (Jollinger, 1979). The main cause of river flooding in the Himalayan Region is the occurrence of heavy rainfall during the monsoon, mostly southern slopes of Churia and Mahabharat receive more rainfall.

In the USA, the most significant floods occur on the Mississippi River, resulting from regional rainfall and snowmelt events that cause slow rise on the rivers and extend for days or weeks. The magnitude of flooding's increases moving downstream from the headwaters to the mouth. Short, intense rainfall events can cause flash floods or quick rise and fall floods on the tributaries but do not normally affect the mainstream. Mississippi River floods and flood damages are significant enough to merit regional or national attention occurred on the Mississippi in 1849, 1850, 1882, 1912, 1913, 1927, 1936 1950, 1973, 1993 and 2001. The largest flood flow recorded on the lower Mississippi was 64,500 m³ s-1. The 1993 flood flow on the upper Mississippi exceeded 28,000 m³/s (Galloway, 2004).

In Kisangani in northern Democratic republic of Congo, the Tshepo and Congo rivers flooded their banks in late 1997 and early 1998, forcing approximately 10,000 people from their homes.in some villages along the rivers greater than 80 percent of homes were destroyed. the flooding was accompanied by an outbreak of cholera that killed hundreds of people (UNDP REPORT)

In the year 2000 Mozambique was hit by the worst flood event in 50yrs, which caused enormous destruction and losses and left a quarter of the population affected with around 540,000 people displaced and one hundred killed. In the following year 2001, the country was flooded again in the central region, as well as in 2003 in the northern region (UNDP report)

In Togo, flooding displaced more than 7,000 people and affected many more in the savannah region of northeastern, which according to recent UN studies is the country's poorest region with "alarming" rates of acute child malnutrition. Floods are reported to have watershed out roads and bridges, completely cutting access to about half of Kpendajal prefecture, which had a population

of an estimated 123,000. other affected are Oti, Tone and Tandjoure (IRIN, All Africa 6 September 2007)

Flooding, discharge of excessive water exceeding the channel capacity, and flash floods, originating from thunderstorms are the most pervasive and costly natural hazards faced by the Nepalese people. The problem of flood hazard is particularly prominent in the southern plains of Nepal where the gradient of the river channel is very low. Apart from hydrological phenomenon, the presence of natural or man-made obstructions in the flood path such as bridge piers, floating debris, weirs, barrage, outburst of landslide dam, glacial lake, failure of different structures, etc. causes flooding. Some of the rivers are dry during the winter season, while in the monsoon they become active and cause immense damage (Dixit, 1999). Therefore, it is not uncommon to have floods in a country having such large rivers and monsoon.

Flood occurrence in one country can affect neighboring countries. For instance, in the Mozambique floods of 2000 was triggered by extremely heavy rain that lasted five weeks in South Africa. Botswana received 75% of its annual rainfall in only three days, and Cyclone Eline hit, bringing more heavy rainfall. The rain from Botswana and other Southern African countries drained into the Limpopo, Zambezi and other rivers which flow through Mozambique to the sea. These rivers eventually burst their banks, causing severe flooding in Mozambique. This resulted to displacement of people, destruction of properties and outbreak of waterborne diseases. The floods were linked to the rapid urbanization in South Africa (BBC, 2000).

Riverine floods are the most dominant floods in Kenya. River floods mostly occur along floodplains or wash lands as a result of exceeded stream flow capacity leading to over spilling of the natural banks or artificial embankments (Smith & Ward, 1998). Riverine floods include extreme events like the El Niño flood of 1997-98 with a low flood probability of 35-40 years (Q 35-40) and the more frequent small floods with a high flood probability of up to five years (Q5). Major rivers in the country such as Nzoia, Nyando, Yala, Athi, Nairobi and Tana experience seasonal river floods emanating from the country's highlands that receive high annual rainfall ranging from 1600-2000mm. These rivers flooded during the 1997/98 rains (Smith & Ward, 1998).

An estimated 723,000 people were affected by devastating floods in kenya in November 2006 according to reports from the relief Web organization. The heavy rains caused rivers to overflow their banks and inundate villages and farmland, produced landslides and increased the risk of cholera and the other diseases that spread when water and sanitation systems were overwhelmed by floods. The united Nations world food program estimates that emergency workers had been unable to reach as many 150,000 affected people due to impassable roads (UN-WFP,http:/col.gsfc.gov).

Heavy rains experienced during the latter part of 2009 resulted in flooding in the coastal regions of Kenya which continued into January 2010 causing flooding incidences in Various parts of the county. Rapid assessments in the affected regions indicated that at least 30 human deaths were reported in various parts of the country, with some 30,000 people in the direct need of emergency aid (Kenya Red cross society, 2010)

At the moment, kenya still lacks a comprehensive flood risk management policy and relies on the existing flood risk reduction measures and strategies proposed in the draft national Disaster management policy of 2007. The draft revised in 2009 proposes a continuum of measures such as preparedness and prevention with the specific strategies such as flood risk vulnerability assessment and mapping an integral flood management (IFM)projects and the strategy for flood management in lake Victoria Basin of 2004 .However ,there are a number of sector specific pieces of legislation that attempt to provide policy direction on flood management in kenya. These include; the water Act, the environmental management in Kenya. These include the water act among others (GOK, 2004, GOK, 2007; WMO) under the chiefs Act (CAP 128), power is vested in the minister to call upon local chiefs to enlist the services of any able-bodied adult to respond to an emergency such as floods (Draft National policy for Disaster management in kenya, Feb 2009 national development process-including through

In 2002, kenya formulated a national policy on Disaster management to institutionalize mechanism for addressing disaster (GOK, 2004b). The policy emphasizes preparedness on the part of the government, communities and other stakeholders in disaster risk reduction activities. It aims to establish and strengthen disaster management institution, partnerships, networking and mainstreaming of disaster risk reduction in the development process so as to improve the resilience of vulnerable groups to cope with potential disasters. It calls for the establishment of a

national Disaster Management Agency (NADIMA)through an act of parliament .The policy integrates disaster risk reduction in the national development process-including through the medium -Term expenditure framework, poverty reduction strategy paper (PRSP), National Development Plan, National poverty eradication plan, economic recovery strategy for wealth and employment creation

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River floods affect both the rural and urban areas in form of flash and urban floods. Flash floods have a characteristic short duration and steep rises and rapid falls of flood levels. They occur abruptly without much warning as a result of an accelerated runoff or sudden dam failure and are quite destructive due to their sudden occurrence. The arid and semi-arid lands of Kenya and urban areas are particularly vulnerable to flash/sheet flooding. Urban floods result from over spilling or surface ponding or when urban storm water drains become surcharged and over flow (Smith & Ward, 1998).

Rainfall intensity, duration and flooding

Rainfall is the most important climatic variable in Kenya that supports most livelihoods, including rain-fed agriculture. An annual rainfall pattern in the country indicates high spatial and temporal variations with a strong bimodal seasonal trend across the country. The months of March-April-May (MAM) and October-November-December (OND) record the highest rainfall peaks. The two rainfall seasons are commonly referred to as the "long" and "short" rains respectively. Torrential rainfall experienced during the wet months often translates into high stream/river flow (runoff) in permanent and intermittent streams/rivers across the country resulting to seasonal floods. (Osbahr and Viner 2006; WRI et al., 2007) indicate that rainfall seasons can be extremely wet and erratic

resulting to both large and small river devastating floods like the El Niño floods of 1997/98 with significant socio-economic impacts.

Flood producing rains in Kenya are often driven by complex climatic variability phenomenon such as the El Niño Southern Oscillation (ENSO)6 and tropical storms. Ogallo (1988; 1989, 1993) cited in ICPAC (2007) notes that specific floods in Kenya have been associated with El Niño. Kenya was amongst the 16 worst affected tropical and Pacific Rim countries during the 1997/98 El Niño (Gadain et al., 2006, Glantz, 2001). The 1997-98 El Niño floods of the century as they have become to known resulted in severe floods after major rivers in the country attained record peaks leading to widespread riparian floods costly impacts. Kenya is one of the countries in the Greater Horn of Africa (GHA) region that has been identified as having a strong ENSO signal (El Niño and La Nina). The warm and cold ENSO cycles are often, but not always associated with above/below average rainfall amounts that lead to floods and droughts respectively.

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Flash floods have a characteristic short duration and steep rises and rapid falls of flood levels. They occur abruptly without much warning as a result of an accelerated runoff or sudden dam failure and are quite destructive due to their sudden occurrence.

The arid and semi-arid lands of Kenya and urban areas are particularly vulnerable to flash/sheet flooding. Urban floods result from over spilling or surface ponding or when urban storm water drains become surcharged and over flow (Smith and Ward, 1998). Urban floods are common in major cities and towns in Kenya including Nairobi, Mombasa, Kisumu, Nakuru, and Garissa.

Many of these urban areas experienced floods during the El Niño rains of 1997-98. The cause of urban floods in Kenya is mainly poorly maintained drainage systems. Urban floods mainly affect residents of informal settlements (slums) mushrooming in the country's major cities. (ICPAC, 2007; UN/OCHA, 2006). The second largest city of Mombasa experienced severe flooding in October 2006 that affected 60,000 people in the city and coastal region (Awuor et al; 2008). The most recent urban floods occurred in Kisumu city between 20-21 September 2009 leaving 150 families displaced and 4 feared dead. This flood was linked to El Niño rains that were forecasted by the Kenya Meteorological Department (KMD) during the short rain season of 2009.

Rainfall is the primary input to most hydrological systems, and a key issue for hydrological science and practice is to assess the importance of the spatial structure of rainfall and its representation for flood runoff generation. At the point when the waters of a flood emerge specifically from precipitation, atmospheric factors can be distinguished as straightforwardly in charge of the occasion (Doswell, 2003). That is, rainfalls happen that are well past the average qualities for the influenced area. It is just when those rainfalls surpass the average that land which is generally dry can be influenced; that is, a flood happens. In this way, the precipitation sums required for floods can't be characterized in total terms (Doswell, 2003). A precipitation occasion that causes a flood in one area may be well inside the limits of what is ordinary for another area. As a rule, the limit for flood-delivering rainfalls increments as the yearly average precipitation for an area increment (Doswell, 2003). Once more, hydrological factors regularly add to a river surge. While singular thunderstorm systems can cause flash floods, river floods are typically the consequence of a dormant concise scale climate pattern. Restricted substantial precipitation occasions happen commonly amid a time of days or even months, each contributing a lot of precipitation to the tributaries, which at that point release into the main stem of a river. The stream rises continuously in light of all the information precipitation. The stream surge capability of a circumstance can be expanded by simultaneous snow melts and different factors other than precipitation (Doswell, 2003).

In Kenya, rainfall is an important climatic variable that supports most livelihoods, including rainfed agriculture. An annual rainfall pattern in the country indicates high spatial and temporal variations with a strong bimodal seasonal trend across the country. The months of March-May and October-December record the highest rainfall peaks. The two rainfall seasons are commonly referred to as the "long" and "short" rains respectively. Torrential rainfall experienced during the wet months often translates into high stream/river flow (runoff) in permanent and intermittent streams/rivers across the country resulting to seasonal floods. Rainfall seasons in Kenya can be extremely wet and erratic resulting to both large and small river devastating floods like the El Niño floods of 1997/98 with significant socio-economic impacts (Osbahr & Viner, 2006; WRI *et al.* 2007).

Rain that results to floods in Kenya are often driven by complex climatic variability phenomenon such as the El Niño Southern Oscillation (ENSO) and tropical storms. Ogallo (1988; 1989, 1993) cited in ICPAC (2007) notes that, specific floods in Kenya have been associated with El Niño. As greenhouse gases increment in the environment, the climate system is warming on the grounds that these gases trap more heat (Hartmann, Klein, Rusticucci, Alexander, Brönnimann, Charabi, Dentener, Dlugokencky, Easterling, Kaplan, Soden, Thorne, Wild and Zhai, 2013). The large water bodies are additionally warming, particularly at the surface, what's more, this is driving higher evaporation rates that, thus, expands the measure of water vapor in the climate. A hotter air can hold more water vapor, driving thus to more exceptional precipitation. The 1°C temperature rise that has just happened, together with expanding evaporation, has prompted an expansion of around 7% in the measure of water vapor in the climate (Hartmann et al. 2013).

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2.3 Budalang'i flood plains

Budalang'i flood plains lies to the north of Lake Victoria close to Kenya-Uganda border's in Busia County. The flood plains are mainly on the mouth of River Nzoia and partly on River Yala (UNEP, 2013). The River Nzoia main streams flows from an elevation of about 2288 a.s.l of the Elgeyo Escarpment and the Cherangani Hills. Its tributaries, which flow from the high slopes of Mount Elgon, attain maximum elevation in the river's basin and are estimated to be about 4300 m a.s.l. (Paolo *et al.* 2013).

The Budalang'i area is prone to flooding, and as a result dykes were constructed between the years 1977-1984 to mitigate the floods. However, with the growing population large areas of forests upstream were cleared for human settlement and crop farming. Consequently, erosion and foreseeable soil slippage and landslides occurs on annual basis due to floods (UNEP, 2009). The accumulation of many years of sediment in the river bed has made the channel of the river course to be above the general level of the floodplain as a result overbank flow across the 400–600-m-wide dykes causes massive flooding. The sediment accumulates and reduces the capacity of river channel so that the river overflows the banks forming the delta (Paolo *et al.* 2013).

2.4 Flood management and control

The management of floods as problems in isolation almost necessarily results in a piecemeal, localized approach. Integrated Flood Management calls for a paradigm shift from the traditional fragmented approach and encourages the efficient use of the resources of the river basin, employing strategies to maintain or augment the productivity of floodplains, while at the same time providing protective measures against the losses due to flooding. Sustainable development through Integrated Water Resources Management aims at the sustained improvement in the living conditions of all citizens in an environment characterized by equity, security and freedom of choice.

Integrated Water Resources Management necessitates the integration both of natural and human systems and of land and water management. Both population growth and economic growth exert considerable pressure on the natural resources of a system. Increased population pressure and enhanced economic activities in floodplains, such as the construction of buildings and infrastructure, further increase the risk of flooding. Floodplains provide excellent, technically easy livelihood opportunities in many cases. In developing countries with primarily agricultural economies, food security is synonymous with livelihood security.

The ecosystem approach is a strategy for the integrated management of land, water and living resources, a strategy that promotes conservation and sustainable use in an equitable manner. Both

Integrated Water Resources Management and Integrated Flood Management encompass the main principles of the ecosystem approach by considering the entire basin ecosystem as a unit and by accounting for the effects of economic interventions in the basin.

Environmental sustainability of the flood management options is one of the prerequisites in IFM. Sustainable and effective management of water resources demands a holistic approach, linking social and economic development with the protection of natural ecosystems and providing appropriate management links between land and water uses. Therefore, water related disasters, such as floods and droughts, because they play an important part in determining sustainable development, need to be integrated into water resources management.

As early as 2000-3000 BC the ancient civilization of Mesopotamia had a system consisting of banks, channels and regulation constructions to ensure safe settlements and good working agriculture (Haggstrom, 1999). Thus, flood control is not a new engineering field but has of course evolved over history in both terms of technical solutions as well as in planning and operation. Flood management is today a large field involving different competencies and many subjects, all with their own difficulties and challenges. The two greatest are how to identify and understand flood risk, and how to handle and reduce that risk by implementation of measures. An integrated approach that combines the whole field is preferred and recommended (Jha et al. 2012).

In Netherlands a methodology applying spatial planning and changes in a wide range of parts of strategy making has been utilized with the end goal that Instead of raising the level of the dykes, dykes have been moved to open up floodplains. Swamps and surge water stockpiling regions have been made to briefly store surge waters when required. Floodplains levels have been brought down to take into account more prominent surge water stockpiling. Homes and families have been migrated to take into consideration extended floodplains (Murray, 2017). North Yorkshire chose to 'moderate the stream' by catchment scale demonstrating to focus on the usage of an extensive variety of measures to diminish, moderate and store surge water in the scene and decrease the size of surge age downstream. Following itemized hydrological displaying, directed measures have been actualized in the correct areas along those tributaries and regions inside every one of the catchments that offer ideal advantage for surge lessening (Murray, 2017). Dublin the legislature brought down segments of parkland by the expulsion of very nearly 60 million liters of soil and the arrangement of built wetlands (large stormwater retention ponds) comprised of five

disconnected lakes, controlled by weirs, were developed on the floodplain. The subsequent brought down territories fill in as a holding territory to hold surge waters and enable them to be gradually discharged once more into the waterway as the surge dies down (Murray, 2017).

Implementing non-structural measures such as operation of flood forecasting and warning system, preparation of hazard maps, and development of evacuation and precaution, improves adaptive capacity of related government agencies and communities.

Development of legal systems such as development regulation for the flood prone areas, development regulation of upstream forest to retain water and farmlands to store flooded water, improve adaptive capacity.

If adequate public information is implemented, and inhabitants are educated and responsive on disaster and risk management issues, they would implement appropriate precautionary measures and actions in times of disaster, hence, improving their adaptive capacity.

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Check dams and diversion channels have been worked by the Ministry of Agriculture to offer assurance to horticultural fields along River Nyando's flood plain. Just 8 km of dykes, 4 km on

either bank were set up in the lower spans of River Nyando continuously 2004. (WMO and GoK, 2004; Otiende, 2009). The Seven Forks Dam venture comprises of a progression of multipurpose reservoirs in the upper catchment of Tana River. Notwithstanding filling in as flood control measures, the dams were produced as a major aspect of the hydropower plot and for water system utilize. The Radio and Internet (RANET) program offers an open door for dispersing flood forecasts to vulnerable rustic networks in the flood inclined territories of Kenya. The administration organizes flood crisis, help and recovery reaction measures through the Rapid-Onset Disaster Committee and District Steering Committees. Philanthropic exercises are embraced in the influenced locale with help from different organizations associated with compassionate reaction, for example, the KRCS, UN offices and NGOs (UN-OCHA, 2006; Otiende, 2009).

Dykes are found along the tremendous system of changeless rivers in Lake Victoria Basin that encounter regular flooding. The dykes on River Nzoia measure 36.4 km long with 16.2 km in the northern and 18.4 km on the southern banks of the River. They offer assurance to inhabitants in Busia and Budalang'i locale in western Kenya (GoK, 2007; Otiende, 2009). The development of River Nzoia dykes started in 1977 and was finished in 1984 at an expense of Ksh 17 million (WMO and GoK, 2004; Otiende, 2009).

2.5 Theoretical framework (Strahler's 1952)

The study adopted Strahler's (1952) stream ordering theory. In Strahler's ordering the main channel is not determined; instead the ordering is based on the hierarchy of tributaries. The ordering follows these rules:

- 1. if the node has no children, its Strahler order is 1.
- 2. if the node has one and only one tributary with Strahler greatest order i, and all other tributaries have order less than i, then the order remains i.
- if the node has two or more tributaries with greatest order i, then the Strahler order of the node is i + 1.

Strahler's stream ordering starts in initial links which assigns order one. It proceeds downstream. At every node it verifies that there are at least 2 equal tributaries with maximum order. If not, it
continues with the highest order; if yes, it increases the node's order by 1 and continues downstream with the new order.

2.6 Conceptual framework

The conceptual framework below (**Figure 1**) presents the relationship between the study variables. The independent variables are: basin morphometry and rainfall characteristics while the dependent variable is flooding in Budalang'i basin. Flood management and control practices usually play a critical role in ensuring the reduction of effects of flood on the inhabitants.



Figure 1: Conceptual framework

Source: Researcher 2018

2.7 Gaps in the literature review

- There is no much documentation information existing on river morphometry and flooding
- River Nzoia basin morphometric characteristics in relation to flooding in Budalangi area has been scantly studied

CHAPTER THREE: STUDY AREA

3.0 Geographical location and introduction to physical features of River Nzoia basin

This chapter covers the location, climate, topography, demography and livelihoods, and soils of the study area. The focus of this study is the section of the Budalang'i that is prone to flooding.

3.1 Location

Budalang'i area is located western part of Kenya near the shores of Lake Victoria. It lies between the latitudes 0^{0} 1' 30" S and 0^{0} 11' 7"N and Longitudes 32^{0} 54' 29"E and 34^{0} 05' 46" River Nzoia navigates the area periodically causing floods in the lower areas as it enters Lake Victoria.

River Nzoia is one of the largest rivers in western Kenya. The main stream of the river flows from the western side of the Elgeyo Escarpment (Sergoi, Sosiani and Kipkelion tributaries) and the Cherangani Hills (Chepkotet and Kaisungur tributaries) from an elevation of approximately 2286 m above mean sea level. Its tributaries, which flow from the high slopes of Mount Elgon, attain maximum elevation in the river's basin and are estimated at about 4300 m above mean sea level. The tributaries in Mount Elgon include Kuywa, Sioso, Ewaso, Rogai and Koitobos.

The total catchment area drained by the basin's river network is about 12,950 km² when measured from Lumbwa Ferry area. Of this about 90.8% (11,667 km²) is composed of a land area, which is relatively flat, with fairly deep soil, continuous vegetative cover, and shows a relatively stable landscape. The mean slope of the basin is 0.071%. The remaining land area is either covered by swamp (5.4%), areas of localized instability, sheet or gully eroded areas or steep sloping areas with rock outcrops.

The average basin elevation is estimated at 1917 km above mean sea level, while the length of the mainstream is 275 km from source to mouth. The river follows the 275-km course with a mean slope of 0.010% from its source to discharge into Lake Victoria at about 1000 m above sea level. It enters the lake a short distance to the north of the Yala Swamp. The upper catchment is a high rainfall zone having a mean annual rainfall of between 1500 and 1700 mm per annum, while the flood plain is a low rainfall zone with mean annual rainfall below 800 mm per year.



Figure 2: Map of the study area

Source: RMCRD, 2018

3.2 Climate

Climate is one of the main factors causing disasters due its abundant precipitation that leads to dam breaks, flooding, soil erosion, water pollution, famine and outbreak of water-related diseases (Odada et al., 2006).

The climate of Budalang'i area ranges from semi humid to semi-arid, semi humid and sub humid characterized by mean annual temperatures of $20-22^{\circ}c$, mean maximum temperature of $26-28^{\circ}c$, mean minimum temperature of $14-16^{\circ}c$ and absolute minimum temperature $6-8^{\circ}c$ (Kenya metrological station). The area is classified as moderately warm in relation to temperature. The area experiences four seasons in a year as a result of inter-tropical convergence zone (ITCZ). There are two rainy seasons and two dry seasons; shorts rain which occurs from October to December) and long Rains (March to May). The local relief and influences of the Lake Victoria modify the weather patterns

3.3 Topography

Topography and size of the basin influences how much and how quickly rainwater reaches the river network. Topographic characteristics that affect runoff response time include watershed shape, drainage pattern, water shed slope and stream channel slope. The topography of a region also determines land-use patterns and the patterns and intensity of erosion and landslides. It influences the gradient of rivers such that in areas with steep slopes, water flows rapidly into river channels. In such cases, the flood stage of a river is reached rapidly. In flat plains, the gentle slope reduces the velocity of flow leading to formation of meanders and flood plains. Water in such areas is not transported rapidly and tends to pile up making the low-lying zones more liable to flooding.

The topography determines land-use patterns and intensity of erosion and landslides. For instance, there is a rapid flow of water in areas with steep slopes thereby causing the flood stage of a river to be reached faster. On the other hand, the velocity of water flow is reduced in flat plains or in areas with gentle slopes. This leads to formation of flood plains and meanders hence pile and become more predisposed to flooding for example Kano Plains (Opere, 2013).

The Altitude of Budalang'i area ranges between 1100-1300 meters above sea level (3600-4300feet). The topography of Budalang'i area is relatively flat, being mostly made up of the alluvial plain of the Nzoia River. The ground generally slopes gently to the south west. A ridge

running north- eastward from Port Victoria prevents the north bank flood plain from draining directly towards Lake Victoria (UNEP, 2006). When the water level in the river Nzoia rises to an extent that it breaks its banks, it is most likely that the larger part of Budalang'i area floods due to the prevailing nature of the land and the soil types



Figure 4: The Terrain of Nzoia River Basin

3.4 Geology

The geologic characteristics of an area are known to influence the drainage patterns, the nature of soils and land-use patterns. For instance, heavy rainfall is more likely to cause flooding in low-lying regions occupied with clay formations because clay formations have low infiltration capacities and therefore surface run-off is generated rapidly. Heavy rainfall, for instance, causes flooding in low-lying areas characterized by clay formations which has low penetration capacities and hence rapid surface run-off (Opere, 2013).

Source: RMCRD 2018

The geology of Budalang'i area is composed of Quartzite, sericite-schist, banded magnetite, quartzite, jaspiilte and greenstone-schist. Altered igneous rocks are still under the microscope and some are of definite pyroclastic origin (R. Murray-Hughes 1933). The rocks along River Nzoia and its tributaries are mostly extensive outcrops of medium grained Kavirondian greywacke, fine grained argillate and nyanzian epidiorite or metabasalt (L.D. Sanders 1965).

3.5 Soils

Different types of soils are found in Busia County. Some parts that of County that experiences flooding is poorly drained due to dominance of clay soil type. In the swampy areas such as the Yala Swamp, have also heavy clay soil making it very difficult for cultivation during the dry or wet seasons, although some other areas have good well drained soil that support cultivation (Kenya Soil Survey, 1991). Generally, the type of soils found in the County are: cambisols, ferrasols, regosols, arenosals, gleysols, vertisols and fluvisols. Gleysols and rrenosols which are found on the Southern side of Nzoia River; cambisols, ferrasols and regosols which are found on the Northern and Southern Sides of Nzoia River (Kenya Soil Survey, 1991).

The soils of the floodplain in the lower reaches of the Nzoia River are all alluvial. The river meanders in the floodplain depositing silt during seasonal floods. The major parts of the floodplain contain black cotton soils, while other areas have coarse textured sand silt mixture. In some places, there exist saline soils. The upper Nzoia Basin contains extensive seasonal swamp areas in the high and medium rainfall zones that are mainly utilized for grazing due to poor drainage. Once the river bursts its banks, the resultant flooding displaces close to 30,000 persons every rainfall season. People are usually evacuated to safer higher grounds during floods. Entire homesteads are swept away, property and crops worth hundreds of thousands of shillings are lost and many people perish as rivers break their banks, rendering large areas of land inaccessible.

The soils in Budalang'i Area comprises of drained to well drained shallow, dark red to brown, friable, sandy clay loam to clay. Some places are characterised with rocky, boulder and stony, with acidic humic topsoil. Most of the area has very poorly drained, very deep dark grey to black firm clay with an acidic humic top soil and in many places peaty. The area has also well drained, moderately deep to very deep, red very friable, sandy clay to clay on the valley sides. This has in

turn attracted farming activies along the river valleys. In case of a flood event, water takes longer time stagnating in most of the area because of the clay soil that tend to retain water for long time. The alluvial soils and annual rainfall of between 750 and 1,015 mm support small-scale agriculture as well as livestock keeping (Bunyala District Report, 2008). The area is a host to the Bunyala Irrigation Scheme, which cuts across the boundaries of Busia and Siaya counties and covers about 1,734 acres.

The basin has the potential for growing a wide variety of crops and producing livestock, due to its good soil and gently sloping terrain. However, the high-water table in the floodplain renders the area uninhabitable during much of the wet season. Population growth and pressure for land however, continues to push people to these flood-prone areas. The search for more agricultural land also spurs encroachment into Nzoia River wetlands. By settling in the flood plain areas, more people and households in Bunyala are at risk of flood hazards (Onywere et. al. 2011; Bunyala District Report, 2008).

3.6 Demography and livelihoods

According to the KNBS (2013), Budalang'i has a total population of about 63,000 people and 15245 households (2013 Census), who inhabit the region mostly affected by floods. Of this, 54% of the total population is made up of women, while 60% is made up of children between 0 and 14 years. The entire Budalang'i has 4 county Assembly Wards with 6 locations and 18 sub-locations. The areas affected by River Nzoia which divide Budalang'i into two sections, suffer a lot during the over flooding of this river since the dykes no longer contain the floods as required. The accumulation of many years of sediment in the river bed has made the channel of the river course to be above the general level of the flood plain as a result overbank flow across the 400–600-m-wide dykes causes massive flooding.

3.7 Drainage

The hydrology of lower Nzoia area lies in Lake Victoria drainage basin and number of significant swampy areas exist in the area including Mundere Swamp to the north of the river and Yala Swamp to the south. Yala Swamp is fed by River Yala and is one of the largest swamps in the area as it drains into Lake Victoria via a number of small channels including Ndekwe stream, which also carries drainage from Bunyala pilot irrigation scheme (UNEP, 2009). The River is about 315km

and drains a total catchment area of 12, 696km². The size of Lower Nzoia basin is 2593km². Along the river the slope reduces from 0.5% in the upper reaches to 0.04% in the lower reaches over the last 30km, which is in the area of the project (Paolo *et al.* 2013).

The total catchment area drained by the basin's river network is about 12,950 km2 when measured from Lumbwa Ferry area. Of this about 90.8% (11,667 km2) is composed of a land area, which is relatively flat, with fairly deep soil, continuous vegetative cover, and shows a relatively stable landscape. The mean slope of the basin is 0.071%. The remaining land area is either covered by swamp (5.4%), areas of localized instability, sheet or gully eroded areas or steep sloping areas with rock outcrops.

The average basin elevation is estimated at 1917 km above mean sea level, while the length of the mainstream is 275 km from source to mouth. The river follows the 275 km course with a mean slope of 0.010% from its source to discharge into Lake Victoria at about 1000 m above sea level. It enters the lake a short distance to the north of the Yala Swamp. The upper catchment is a high rainfall zone having a mean annual rainfall of between 1500 and 1700 mm per annum, while the flood plain is a low rainfall zone with mean annual rainfall below 800 mm per year.

The width of the channel decreases gradually from about 50 m at 140 km inland to around 40 m in the upper reaches of Kakamega, Bungoma, Trans Nzoia and Uasin Gishu Districts where the altitude is high, and the slopes are steep. The sediment accumulates and reduces the capacity of river channel so that the river overflows the banks forming the delta. The flow regime of the Nzoia is varied and is occasionally as low as 20 m³/s and with extreme floods that may surpass 1100 m³/s, which is the proposed protection level for the dykes for a 25-year, return flood. Siltation is heavy. Earlier estimates of siltation rates by Italconsult are in the order of 158 tonnes per day; however, recent assessments by Lake Basin Development Authority (LBDA) for the period 2000–2003 put it at 574 tonnes per day, which is thrice the value. The discharge varies from a low flow of 2.8 m3/s to a 100-year flood flow of 930 m³/s.



Figure 5: The River Nzoia watershed and the main water catchment areas of the basin

Source: RMCRD 2018

CHAPTER FOUR: RESEARCH METHODOLOGY

4.0 Introduction

This chapter presents the research methodology applied in conducting the study. It covers the research design, target population, sampling design, data collection and data analysis techniques.

4.1 Research design

The study adopted a descriptive research design. Descriptive research collects data and makes deductions about a population of interest at one point in time. A descriptive research design describes people responses to questions about a phenomenon or situation with aim of understanding respondent's perceptions from which truism is constructed (Kim, 2009). The aim is to gather data without manipulation of research context and deals with natural phenomena where the researcher has no control over variables (Mugenda & Mugenda, 1999).

4.2 Target population

According to the KNBS (2013), Budalang'i has a total population of about 63,000 people and 15245 households (2013 Census), who inhabit the region mostly affected by floods. The target population consisted of 237 households which were identified as living in the flood risk areas.

4.3 Sample design

4.3.1 Sample size

Using the following formula (Nassiuma, 2002)

Sample size (n) = $\{NC_v^2\}/\{C_v^2+(N-1)e^2\}$

Where:

Cv = coefficient of variation (0.5)

e = tolerance at desired level of confidence (0.05) at 95% confidence level.

N = target population

Therefore: n =. $\{237*0.5^2\}/\{0.5^2+(236*0.05^2)\}$

=71 households

Households were the unit of analysis because of the governmental structure in the study area because they represent the area with identical inhabitants with respect to population physiognomies, economic status, and living conditions.

4.3.2 Sampling procedure

Simple random sampling was used to select the households to make sure that they provide a true sample of the entire population and allow each household an equal chance of being selected. This also ensured that accurate results are obtained for purposes of generalizing and inferences. A small sample size of 70 households was used for the questionnaire administration.

4.3 Data collection

Primary data on human activities, distribution and seasonal variation of the river was obtained by field observations and photography. Structure questionnaires were used to gather information on flood risk management and flooding occurrences in the study area.

Secondary data was obtained on land use and morphometry from satellite images from Department of Resource Surveys and Remote Sensing (DRSRS), rainfall intensity and duration from the Meteorological Department and soil characteristics from the Soil Survey of Kenya.

Topographical maps of River Nzoia Basin on scale 1:50,000, published by survey of Kenya, obtained from their offices were used as fundamental data source for computing morphometric parameters. These maps are highly endorsed for a study of this nature (Ajibade et. al, 2010). The hard copies of topographic maps of the study area were scanned and saved in GIS compatible format; it was later exported to the ArcGIS 10.2 user interface where the maps were georeferenced and brought to a common spatial reference system for better analysis.

4.4 Data analysis

The major datasets that were used for this study include: Landsat Satellite Imageries, Digital Elevation Data (DEM) and Auxiliary data that included Roads, administrative boundaries, ground truth data for land use land cover classes and topographic map of the area.

1) **DEM**

The Shuttle Radar Topography Mission (SRTM) elevation model was obtained from Regional Center for mapping of Resources for Development (RCMRD) at a resolution of 30 meters. The DEM was used in a Geographical Information System to generate the catchment area and drainage network. Spatial analysis using ArcGIS hydrological tools was carried out to obtain stream orders and to calculate the stream number and stream length.

2) Landsat data processing

The common way to map land use land cover of an area is to use Satellite image classification. For this study, cloud free Landsat images were obtained from RCMRD (table 2) courtesy of United States Geological (USGS) to perform land cover classification.

Date	ID	Sensor	% Cloud
(mm/dd/year)			cover
09242018	LC08_L1TP_17006020180924_01_T1	LANDSAT_8"	0.04
		"OLI_TIRS"	
10021986	LT05_L1TP_170060_19861002_01_T1	LANDSAT_5	2.00
		ТМ	
10081996	LT05_L1TP_170060_19961008_01_T1	LANDSAT_5	1.00
		ТМ	
10232011	LT05_L1TP_170060_20111023_01_T1	LANDSAT_5	12.00
		ТМ	

Table 2: Landsat imagery

Landsat scenes are provided to the users as level 1 under the Geotiff format with a spatial resolution of 30metres. Prior to change detection, the images were preprocessed in Idrisi selva; sub setting the imagery based on the area of interest and image enhancement through combination of various

bands. All satellite data were analyzed by assigning per pixel signatures and differentiating the land cover into four classes based on specific Digital Number (DN) value of different landscape elements. The delineated classes were Farmlands and bare areas, marshy areas, Forest, and Water bodies (**Table 1**)

The collected questionnaire data was cleaned, validated, and edited for accuracy, uniformity, consistency and completeness. Microsoft Excel was used to analyze the questionnaires and rainfall data. The main methods used in analyzing quantitative data on rainfall intensity and duration was descriptive statistics using tables and figures. Land use and basin morphometry was assessed from satellite imageries to classify their patterns (Lillesand, Kiefer and Chipman, 2004). Basin morphometry was analyzed using digital elevation models. Assessment and description of flooding and its effects was done from field observations and satellite images.

4.4.1 Stream numbers (sn)

The numbers of streams of different orders in a given drainage basin was counted and classified as per Horton's law of drainage as 1st order, 2nd order, 3rd order and 4th order.

4.4.2 Stream length (sl)

The stream length ratio (RI) was calculated as:

Equation 1: how to calculate stream length

$$(\mathbf{RI}) = \frac{Sn}{Sn-1}$$
$$(\mathbf{RI}) = \frac{Sn-1}{Sn}$$

Where Sn = length of an order, and Sn-1 = length in the next higher order.

4.4.3 Bifurcation ratio (Rb)

Bifurcation Ratio is the number of streams in a low order to the number of streams in the next high order (Horton, 1945). Bifurcation means dividing in two, and can be calculated as:

Equation 2-how to calculate the Burfacation ratio

 $Rb = \frac{number \text{ of stream segment in a given order}}{Number \text{ of streams segments in the next highest order}}$

4.4.4 Relief Ratio (Rh)

Equation 3: how to calculate relief Ratio

Relative Relief = <u>Maximum basin relief</u> Maximum basin length

4.4.5 Drainage density (Dd)

Equation 4: how to calculate drainage density(Dd)

D = ratio of total channel-segment lengthsBasin area

CHAPTER FIVE: RESULTS AND DISCUSSIONS

5.0 Results

5.1. Rainfall intensity and duration and flooding

Although Budalang'i Sub County receives bimodal annual rainfall of 1,020 and 1, 2700 mm, it has been affected by floods due to heavy rainfall received in the river's upper zones of Trans Nzoia. The rainfall patterns of long rains experienced in upper catchment areas of Cherengani where the river is in it youthful stage and short rains in areas of Bungoma and Kakamega where the river is in its middle stage results in increased volumes of water available within a limited river channel. High volumes in upper catchment areas during heavy rains of long rain seasons are at higher velocity due to sudden change of river gradient. This result in higher volumes received at the lower catchment areas known as flood plains (in-flow) against that volume expected to be released to the lake (out flow). This volume difference causes the river to break its normal banks in the lower zones, also referred to as the flood plains. Since 1961/62 these floods have caused havoc on both Social, psychological and Economic lives among the residents of Budalang'i sub county, Busia County.

The rainfall stations used for this analysis are 8935181 (Eldoret Meteorological station), 8935115 (Eldoret international Airport), 8934096 (Kakamega Meteorological station) and 8834098 (Kitale Meteorological station) since they were readily available and are in the catchment areas of River Nzoia.



Figure 6:Rainfall in 1974, 1980, 1986 and 1992

Source: Field data 2018

In the areas forming the River Nzoia watershed where rainfall data was used shows that rainfall in the years 1974-1992 was experienced throughout the year with peaks from March all the way to October. There have been variations in the various months as the years pass by owing to the climate changes experience in the continent.



Figure 7: Rainfall in 1998, 2004, 2010 and 2014

Source: Field data 2018

The years 1998-2014 the data shows a shift in the rainfall amounts where the intensity increased significantly and peaking in March, April, July, August and September. According to the residents most surges had a tendency to happen in August, which harmonizes with wetter months of the year (July, August, September, April and March). Surges are additionally more exceedingly identified with the aggregate precipitation happening in a spell of rain, than to intensity. There are increased flooding rates in the later years and this can be tied to the increasing rainfall in the recent years. The duration of rainfall in figures 5 and 6 above is all through the year with various peak months. This affects flooding in several ways: it increases run off since it influences the infiltration and ground seepage rates especially by saturating the soils (*Vigliane & Plaschl 2000*). The smallest

ground seepage rates especially by saturating the soils (Viglione & Bloschl, 2009). The smallest long duration rainfall results in relatively large floods because saturation of the catchment is reached. In short duration rains, the catchment does not contribute fully to runoff generation, so the storms must be extremely severe to produce big floods.



Figure 8:Long-term mean annual rainfall, Nzoia Basin

Source: RCMRD 2008

5.2 Morphometric characteristics

Basin morphometry denotes the topographic representation of land by way of area slope, shape, length etc. These factors affect catchment streamflow model through their influence on concentration time. Strahler (1964) groups morphometric relationships into three: (1) linear morphometric relationships, e.g., bifurcation ratio, length ratio, etc (2) areal morphometric relationships, e.g., law of stream areas, drainage density etc and (3) relief morphometric indices e.g. maximum relief, relief ratio, ruggedness number etc.

Geomorphologic features of many drainage basins and their relationship with runoff response in Kenya are inadequately described. However, a number of studies on geomorphological implications on hydrology have been studied elsewhere (e.g. Nyadawa and Muiruri (2001); Patton; (1988), Iturbe, (1982). Patton, working with various basins in North America developed basin

specific regression formulae for predicting floods based on ruggedness number, relief ratio, first order channel frequency and basin magnitudes. Nyadawa and Muiruri (2009) explained the discrepancy of runoff response in two sub basins in Upper Athi in reference differences in bifurcation ratios and form factors but like other previous studies elsewhere the relationships are non-transferrable.



Figure 9: River Nzoia stream order and flood prone areas map Source: RCMRD 2018

N/B-The small dot represents areas less affected by floods while the big dots represents areas that have been intensively affected by floods

5.2.1 Stream order, stream number, stream Length and Bifurcation ratio

Stream order	Stream Number	Stream length	Bifurcation ratio
1	6556	6184.74	2.25
2	2916	3136.44	1.76
3	1660	1672.13	1.90
4	874	777.99	1.47
5	594	455.30	3.99

Table 3: Stream order, number, length and bifurcation ratio

Source: RCMRD 2018

Areas in stream orders 1 and 5 have the highest flood risk because their bifurcation ratio is above 2. The areas traversed by stream orders 2, 3 and 4 have a lower bifurcation ratio thus lower chances of flooding.

Stream Order (Nu)- The first step in drainage basin analysis is to determine the stream orders. Stream order has been defined as a measure of the position of a stream in the hierarchy of tributaries. Hierarchical ordering of streams is necessary to understand the hydrodynamic character of a drainage basin. In this study, the various stream segments of the drainage basin have been ranked according to Strahler's stream ordering system. According to Strahler's law, the smallest fingertip tributaries are designated as first stream order. Where two first stream order meets, a second stream order is formed; where two second stream order meets, third stream order is formed and so on. River Nzoia is a sixth order drainage basin. The total numbers of 12,600 streams were identified as shown in the table above.

Stream Length (Lu)-The total length of individual stream segments of each order is the stream length of that order. Stream length measures the average (or mean) length of a stream in each order and is calculated by dividing the total length of all streams in an order by the number of streams in that order. The numbers of streams of various orders in the basin were queried and their lengths also determined and summed up in ArcGIS 10.5 environment. The stream length in each order increases exponentially with increasing stream order as shown in the table (2) above.

Stream length is one of the most important hydrological features of the basin as it reveals the surface runoff characteristics. Streams of relatively smaller lengths indicate that the area is with high slope and finer textures. Streams with longer lengths are generally indicative of flatter surface

with low gradients. Stream length is a revelation of the chronological developments of the stream segments including interlude of tectonic disturbances. Usually, the total length of stream segments is highest in first stream orders and decreases as the stream order increases.

Bifurcation ratio (Rb)

Bifurcation ratio is related to the branching pattern of a drainage network and is defined as the ratio between the total numbers of stream segments of one order to that of the next higher order in a drainage basin

The computation of bifurcation ratio shows that the values ranged from 1.5 and 3.4 which are within Strahler's range. The average value of Bifurcation ratio is 2.2 which is closer to the lower value of 2. Based on bifurcation ratio, River Nzoia basin flows may encounter delayed time to peak and this is a good characteristic for planning and conveying flood forecasts



Figure 10: River Nzoia Basin DEM Source: RCMRD 2018

5.2.2 Drainage density

Sub basin code	Total area	Area (10 ²)	Drainage density (10 ⁻¹⁾
	length (10)		
BA	89.47	628.14	0.14
BB	105.34	610.32	0.17
BC	142.30	767.49	0.19
BD	110.88	070.22	0.17
BE	181.32	1007.89	0.18
BF	174.04	404.34	0.19
СА	158.51	302.93	0.11
СВ	74.54	02.93	0.13
CC	43.35	07.41	0.20
CD	101.40	520.11	0.20
CE	54.50	229.55	0.24
DA	134.37	518.15	0.36
DB	254.07	024.24	0.37
DC	105.28	371.34	0.45
DE	82.70	354.72	0.23
EA	125.87	427.43	0.29
EB	138.07	378.75	0.37
EC	75.04	240.10	0.30
ED	30.03	128.18	0.24
EE	07.38	399.37	0.17
EF	37.14	383.85	0.27
EG	143.14	542.15	0.20
Mean drainage			0.24
density			

Table 4: Computation of sub-basin drainage densities

Source: Kenya Rivers shapefiles



Figure 11: Spatial distribution of drainage densities within Nzoia basin

Source: RCMRD 2018

Analysis of spatial distribution of drainage density has shown consistent conveyance efficacy with a mean value of 0.24. The value has no control values to base judgement. Based on that, the interpretation is limited to spatial variability and frequent occurrence of floods is attributed to effect of relief ratio.

5.2.3 Relief ratio

Table 5: Relief ratio computation

				Stream	n order n	0		
Sub- Basin	Distance(L)	Height H(M)	Relief Ratio®	1	2	3	4	
BA	61800	1398	0.0226	4	1	0	0	
BB	47560	1386	0.0291	1	1	0	0	
BC	39120	1343	0.0343	20	4	2	1	
BD	49930	542	0.0109	2	1	0	0	
BE	59450	2127	0.0357	11	2	1	0	
BF	61380	2402	0.0391	22	5	2	1	
BH	64960	2542	0.0391	19	3	1	0	
CA	59940	1265	0.0211	1	0	0	0	
СВ	71490	1410	0.0197	3	1	0	0	
CC	48480	900	0.0186	11	5	1	0	
DB	70180	2893	0.04122	13	3	1	0	
DD	43650	402	0.0092	2	1	0	0	
EA	24960	636	0.0255	15	5	2	1	
EB	33450	392	0.0117	6	2	1	0	
EC	38720	380	0.0098	8	3	1	0	
EG	48780	342	0.007	16	9	4	2	

Source: Kenya rivers Shapefiles

Relief ratio is a non-dimensionless parameter which measures the steepness of a watershed. In this case, it is observed that all sub basins whose divide line pass through upper mountain ranges e.g. DB, BH, and BE have high relief ratios of above 0.03and this drops drastically to a low value of about 0.01 for sub basins at the lowland.

The average relief ratios of all the sub basins of similar order shows that there is negative gradient from order 1 and 2 which corresponds with normal trends in basins of uniform relief at the divide, given that basin length increases in magnitude with increase in basin order. This theory of the uniformity of relief at the divide does not hold beyond order 2. This leads to the conclusion that there is no consistency between relief ratio and basin order in Nzoia and no relief at the divide.

5.3 Characteristics of River Nzoia in relation to flooding in Budalang'i

5.3.1. River gradient and increased meanders

The length of the main stream is about 252 km with a fall of about 1200 m (BUCODEV 2017). over this stretch; the river meanders and causes deposition of silt due to the low gradients. The sediment accumulates and reduces the discharge capacity of the river channel so that it over flows its banks causing flooding in lower parts of the catchment.



Figure 12: Reduced river depth

Source: Field data 2018

Given that the upper catchment area receives more rains than the lower parts of the catchment areas coupled with the fact that the change in riverbed gradient is so sudden, there is intense erosion

from the early stages of the river downstream. There is increased deposition on the old river stage at lower catchment basin which results in reduced river depth. This contributes heavily to flooding as high volumes surpasses the river banks and dykes.

5.3.2 River Width

The width of the channel decreases gradually from about 50 m at 140 km inland to around 40 m in the upper reaches of Kakamega, Bungoma, Trans Nzoia and Uasin Gishu counties where the altitude is high, and the slopes are steep. The sediment accumulates and reduces the capacity of river channel so that the river overflows the banks forming the delta. This reduced width of the channel affects the river's capacity to contain high inflows from the upper catchment reaches causing flooding.

5.3.3 Reduced River Velocity in Lower Catchment Areas

A sudden change in river gradient has caused a reduced river velocity. The reduced velocity increases chances of deposition which consequently leads to the accumulation of many years of sediment in the river bed has made the channel of the river course to be above the general level of the flood plain as a result overbank flow across the 600–700-m-wide dykes causes massive flooding. In A nutshell, the main factors contributing to increased incidence of floods especially in the lower reaches of the river Nzoia are:

- 1. Reduced flood carrying capacity of the rivers due to excessive siltation of their bed.
- 2. Settlement of the dykes in some places reducing the effective height, thereby rendering the dyke susceptible to over topping during floods.
- 3. Erosion of dykes due to river attacks making them susceptible to breaches even during floods of lower magnitudes than the design flood.
- 4. Increase in flood discharges due to severe degradation of watersheds cause uncontrolled and unregulated human activity, especially large-scale deforestation.
- 5. There is no programme / provision for routine maintenance of dykes.
- 6. The dykes are, at present, overgrown with trees and shrubs resulting in loosening and cracking of the dykes.
- 7. Excessive encroachment of the flood plains by the people for agriculture and livestock farming and fishery.

5.3.4 Land use land cover changes

Remote sensing data was used to obtain Land use land cover changes. The landcover land use was obtained by detecting changes through classification of selected multitemporal Landsat images between 1986 to 2018. Land use landcover changes were analyzed and calculated in square kilometres and percentage changes.

The possible changes in the land use/cover include deforestation (afforestation), intensification of agriculture, drainage of wetlands and urbanization. Deforestation, which has converse effects to afforestation, affects significantly the characteristics of the stream flow (Calder, 1992). Though considered a myth or folklore (McCulloch and Robbinson, 1993, Calder, 1998) forests are thought to generate rain, regulate low flows, reduce floods, ameliorate soil erosion and sterilize water.

The intensification of agriculture affects the runoff generated through the alteration of evaporation and the timing of runoff. These effects are compounded by the replacement of certain crops, which alters the leaf area index (Calder, 1992). Wetlands do not or only marginally affect the basin's seasonal water balance (Calder, 1998). However, due to the presence of a free water surface and the lack of water stress, the wetland vegetation normally has a high evaporation rate compared to other land covers. This in turn affects the annual stream flow, which is likely to be less compared to other land uses (Calder, 1992). The earth's climate is also changing gradually. In East Africa for example, catchments are displaying a small increase in annual precipitation received and this makes them wetter. These changes affect the quantity of stream flow.





Figure 13: LULC 1986

Table 6: Land use Land cover for 1986

LULC Category	1986	% COVER
Vegetation	64.58	20.86
Farmlands and Bare areas	55.29	17.86
Marshy areas	58.26	18.82
Water body	131.49	42.47
	309.62	100.00

Source: Researcher 2018

Most part of the constituency is covered by the Lake Victoria. The dominant land cover in Budalangi constituency in the year 1986 was Marshy other vegetation followed by marshy areas which occupied 21 % and 19% respectively. Bare areas and farmlands occupied 18%.



Figure 14: LULC 1996

Table 7: Land use Land cover for 1996

LULC Category	1996	% COVER
Vegetation	52.98	17.11
Farmlands and Bare areas	64.73	20.91
Marshy areas	62.23	20.10
Water body	129.68	41.88
	309.62	100.00

Source: Researcher 2018

In the year 1996 Bare areas and farmlands occupied 21%, marshy areas 20% and vegetation 17% of the total area.



Figure 15: LULC 2011

Table 8: Land use Land cover for 2011

LULC Category	2011	% COVER
Vegetation	55.08	17.79
Farmlands and Bare areas	79.09	25.54
Marshy areas	47.19	15.24
Water body	128.26	41.42
	309.62	100.00

Source: Researcher 2018

In 2011 Bare areas and farmlands were the most dominant land use/ land cover occupying 26%. Vegetation and marshy areas had 18% and 15% respectively.



Figure 16: LULC 2018

LULC Category	2018	% COVER
Vegetation	39.23	12.67
Farmlands and Bare areas	99.42	32.11
Marshy areas	42.98	13.88
Water body	127.99	41.34
	309.62	100.00

Table 9: Land use Land cover for 2018

Source: Researcher 2018

The dominant land cover in Budalangi constituency in 2018 is Bare areas and farmlands

occupying 32%. Marshy areas and Vegetation occupy 14% and 13%.

	%	%			%	%			%	%		
LULC	COVER	COVER	%CHANGE	Type of	COVER	COVER	%CHANGE	Type of	COVER	COVER	%CHANGE	Type of
Category	1986	1996	1986-1996	change	1996	2011	1996-2011	change	2011	2018	2011-2018	change
				Positive				Positive				Negative
Vegetation	14.40	17.11	2.71	(increase)	17.11	17.79	0.68	(increase)	17.79	8.79	-9	(Decrease)
Farmlands and				Negative				Positive				Positive
Bare areas	24.32	20.91	-3.41	(Decrease)	20.91	25.54	4.63	(increase)	25.54	32.11	6.57	(increase)
				Positive				Negative				Positive
Marshy areas	18.82	20.10	1.28	(increase)	20.10	15.24	-4.86	(Decrease)	15.24	17.11	1.87	(increase)
				Negative				Negative				Positive
Water body	42.47	41.88	-0.59	(Decrease)	41.88	41.42	-0.46	(Decrease)	41.42	41.98	0.56	(increase)

Table 10: Summary for land use land cover for 1986, 1996, 2011 and 2018

Source: Researcher 2018

LULC PERCENTAGE COVER					
LULC Category	1986	1996	2011	2018	
Vegetation	20.86	17.11	17.79	12.67	
Farmlands and Bare areas	17.86	20.91	25.54	32.11	
Marshy areas	18.82	20.10	15.24	13.88	
Water body	42.47	41.88	41.42	41.34	
	100.00	100.00	100.00	100.00	

Table 11: Land use land cover percentage cover from 1986 to 2018

Source: Researcher 2018

5.3.4.2 Land use/ Land cover change

			CHANGE 1986-	
LULC Category	1986	2018	1996	Type of change
Vegetation	14.40	8.79	-5.61	Negative (Decrease)
Farmlands and Bare				
areas	24.32	32.11	7.79	Positive (increase)
Marshy areas	18.82	17.11	-1.71	Negative (Decrease)
Water body	42.47	41.98	-0.49	Negative (Decrease)

Table 12: Land use land cover changes between 1986 and 2018

Source: Researcher 2018

Remote sensing data showed depicted pressure on the land resources that is seen from the status of land use and land cover change. The intensification of agriculture is seen from the land being converted to medium scale and larger scale sugarcane growing from 84 km² in 1986 through 102.28 km² in 1996 to 262.75 km² in 2011 most of which is in reclaimed wetland areas. Small scale mixed farming increased from 36.29 Km² in 1986 to 244.036 km² in 2005. Trans-Nzoia Plateau occupies north of Nandi escarpment and is heavily cultivated showing a dominance of red iron-rich soils. Existing forest cover is still under threat as seen from clear cut into forest zones

Forest depletion and encroachment to the Nandi Escarpment's hilly area has led to severe erosion as reflected in the high turbidity of the water and landslides often leading to loss of human lives.

The increase in size of the bare ground area is an indication of wetland degradation and a consequent replacement of wetland vegetation with bare soil that is prone to erosion. This contributes to the siltation of the lake thus increasing the lake's water turbidity because the wetland has lost its ability to filter off sediments in water entering the lake. From the satellite images, it was evident that there were levees naturally developed along the riverbank for containment of the river water. The reduction in marshland vegetation perhaps is attributed to the increased intensification of dry land rice cultivation Analysis of the year 2005 images shows that cultivation takes places in the wetland and there is complete drainage of wetland resulting in settlement and cultivation extending to the levees and the flood plain.

The increase of the area under farming, the stream flow increases during the rainy seasons and reduces during the dry seasons, whereas when the area under forest cover is increased the peak stream flow reduces, but when the forest cover is reduced to almost nil there is an increased peak and mean stream flow in the basin.

It is therefore worth perceiving that a decrease in surface runoff would be desirable, as this would also decrease the devastating effects of floods; the rapid expansion of urban centres in the lower parts of the catchment (Rwambwa, Kakamega, and Bungoma) can be said to be a major contributing factor to the annual devastating floods.

The area under vegetation cover decreased between 1986,1996 and 2018 by 9.0% in Budalangi area. But between the 2011 there was an increase in area under forest cover by 0.68%. Agricultural land use showed an increase in areal coverage from 1986 to 2018 by 11.17%. Marshy areas decreased since 1996 to 2018. This could be because of an expansion in riverine farming

In summary, some of the existing land use practices such as deforestation and urbanization enhance flooding. For instance, increased urbanization and deforestation result in reduced infiltration with increased run off in built and urban areas leading to flooding in the lower parts of the Nzoia basin.

5.4 Questionnaire data review

5.4.1 Household characteristics

Out of the 49 respondents, 59.18% were male and 40.82% were female. There was fair representation of the gender in the study, and therefore the opinions gathered were not largely skewed to either side. It also indicated that both men and women were willing to participate in the

study. According to Chanthy and Samchan (2014), floods affects mostly women and yet their concerns are least likely to be addressed. Having a representation from both men and women helps to capture the all interests and make meaningful decision about flood risks.

Large proportion of the respondents were married (44), while 3 were widowed and the rest were separated (2%) and single (2%). Each household had an average of 6 people, with most of the families having lived in the area for an average of 26.67 years. This implies that the respondents have experienced flood for many years and hence their opinions about flood improves the validity of this study.

Marital status	Percentage
Married	89.80%
Separated	2.04%
Single	2.04%
Widowed	6.12%
Grand Total	100.00%

Table 13: Marital status

Source: Field data 2018

5.4.2 Flooding occurrence, risks and impacts on households

A large proportion (93.88%) of the respondent affirmed that they had experienced flood, with only 6.12% stating otherwise. On average, a household lost crops worth KES 15,143, livestock equivalent to KES 8,294, and house damage valued at KES 4,449 per year due to floods. The United Nation Office for Disaster Risk Reduction study in 2015 estimated that the annual loss due to flooding in the Nzoia to be US \$ 1.8 million, and displacement of 12,000 people. The respondents when asked the last time they experienced floods, majority 81.84% said August and April 2016. The rest of the years had less than 11% scores (Figure 5.1). The majority (69.39%) of respondents were of the opinion that floods were likely to occur again within 1 year, with 30.61 % attributing the next flood occurrence on the rainfall intensity. According to Flood Mitigation Strategy (2009), torrential rainfall has been the major cause of floods in Kenya and for Nzoia River, perennial floods are experienced in its lower reaches affecting areas such as Budalang'i and Kano flood plains.
Although, many respondents were aware that the floods would occur again the following year, many (40.82%) have not put in place any measure to safeguard themselves from floods. However, 38.78% of the respondents preferred to move to safer places during the flooding seasons (Table 5.1). The lower reaches of Nzoia River and its tributaries are a source of water for arable farming. And this could be reason as to why most households would prefer to settle in flood areas and move to higher ground when floods are deemed to be severe.



Figure 17: Respondents perception on flood occurrence

Source: Field data 2018

Measures in preparation for floods	Percentages
Bought a boat to help him relocate during flooding	2.04%
Constructed dykes Move to safer areas	2.04%
Constructed water proof wall	2.04%
Maintenance of dykes Move to safer areas	2.04%
Maintenance of dykes	4.08%
Move to safer areas	38.78%
No measures in places	40.82%
Placed stones in the compound to avoid sinking when relocating	2.04%
Planted crops that will take short period to mature	2.04%
Repair dykes by putting a lot of soil on the dykes	2.04%
Repair dykes, Farm in the irrigation scheme	2.04%
Total	100.00%

 Table 14: Flood mitigation measures preferred by respondents

Source: Field data 2018

About halve of the respondents (51.02%) agreed that they get information from the government about floods, while 36.73% Disagreed (Figure 13). Although, 46.94% were of the opinion that the government was not doing enough to end the problem of flooding. Almost all the respondents (97.96%) said they worry often about floods and its related impacts, with 93.88% of the opinion that flood is a treat to their family safety.



Figure 18: Respondent's opinion on availability of information on flood from the government Source: Field data 2018

5.4.3 Factors causing floods

A large proportion (91.84%) of the households were of contrary opinion that deforestation and settlement were not the responsible for the severe effect of flooding along River Nzoia. Instead many attributed the flooding to lack of protective dykes (98.33%); lack of repair of the existing dykes (95.92%) and rainfall intensity (79.59%). In addition, most respondents (95.29%) thought that sand harvesting from the river and change in the river characteristics (67.35%) had no effects on the flooding of River Nzoia.

5.4.4 Flooding risks in Budalang'i

According to Frans (2009), flood risk can be defined in two ways namely;

- Flood risk = (flood) hazard * (exposure) * vulnerability (of the society/ area)
- Flood risk = probability (of the flood) * consequences

Perennial flooding in Budalang'i has had devastating impacts of floods on households in Budalang'i including loss of human life, damage to farms, loss of crops, destruction of property (e.g. Buildings and pit latrines) and disruption of social and economic activities. The following were the flooding risk that were identified in Budalangi floodplains:

5.4.4.1 Displacement of households

Constant ravaging floods in lower parts of the Nzoia basin (Budalang'i) during heavy downpour of rain in up stream of River Nzoia results in displacement of people. Floods of 2012 submerged a relatively considerable surface on either side of the dykes. Households from Bunyala central and North relocated to southern and Northern dykes respectively.



Figure 19: Flooded households

Source: Field data 2018

5.4.4.2 Damage to Farmlands

Considering narrow river channel and numerous meanders of River Nzoia, high water volumes from upstream of the basin causes water overflow on banks resulting in a fair coverage of surface runoff. Farms within and outside dykes are washed away causing a ripple effect of hunger and consequent increase in price of food stuffs.



Figure 20: Households along river banks

Source: Field data 2018

5.4.4.3 Damage to Bridges and Roads

Owing to the floods of 2007, Ndekwe bridge which connected Mabinju and Rukala sublocations of Bunyala sub county was swept away rendering locals helpless as far as transportation and movement was concerned. Sections of roads within the area were destroyed eg. Rukala, magombe Igigo and pathways along the dykes were also washed away disrupting road communication in aforesaid areas.

5.4.4.4 Disruption of learning programmes

Frequent flooding in Bunyala has caused undesirable effects on learning programmes in many of schools in the area. Floods of 2008 ravaged Makunda Secondary school during the time when candidates were seating their national examination. The school had to take candidates to the neighbouring school-St, Anne Bunyala girls to write their final exams.

5.4.4.5 Economic

During floods (especially flash floods), roads bridges, farms houses and automobiles are destroyed. People become homeless. Additionally, resettlement of the displaced people come at a cost and takes longer enough for business to normalise.

5.4.4.6 Environment

The environment is equally affected when chemicals and other hazardous substances end up in in the water and eventually contaminate the water bodies that floods end up in. water hyacinth have evaded the mouth of river Nzoia as it finds its way into the lake. These have caused blocking of the channels of the major river and its tributaries including Ndekwe stream.

5.4.4.7 People and animals

Many people and animals die during floods as some are injured and others made homeless. Electricity supply is partly affected in Budalang'i and its environs causing disruption of businesses. Flooding too comes with disease such as malaria, typhoid, and dysentery among others.

5.5 Flood mitigation and control measures

5.5.1 Community non-structural measures for flood control

The other non-structural cause of flooding in Budalang'i is Poor on-farm activities including wrong farming methods on areas closer to the main river channel and dykes has caused surface erosion which consequently result in opened up flood plain making it vulnerable to flooding.

Cultivation on steep slopes without applying soil conservation measures promotes soil erosion and rapid generation of surface runoff.

5.5.2 Sensitization activities would remain necessary

Early warning systems plays an important role in reducing flood damage. An attempt to use hazard maps is erected at strategic sites on roads and walkways. However, these have the limitation of being unsustainable as maintenance at these sites appeared costly to the communities at the villages. Communities that have been given some evacuation drills and are informed on early warning

5.5.3 Methods shall be developed.

Investment in structural flood control measures such as dams and dykes. This was achieved through public/private investment in multipurpose dams for water harvesting, storage and production of electricity. This is exploited for fish farming and small-scale irrigation projects that target local communities in the study area; and conduct floodplain mapping, land use planning, and enforcement of laws on land use, agriculture, and settlement. For example, homesteads should not be constructed in the flood plains, near dykes, or along river banks.

Efforts are directed on training in disaster management, emergency communication and establishment of disaster management plan. Such disaster management plans would best be implemented through establishment and development of community-based flood management organization.

5.5.4 Structural measures for flood control

Structural measures in flood management basically emphasize river bank protection and harnessing for use, the flood waters. Damaged river banks are being rehabilitated. The drainage systems are widened and desilted by the community. Maintenance of water canals and irrigation channels has proven to be costly for government institutions such as the National Irrigation board. Construction of dams on upper parts of the stream to harness river water at higher velocities for other socio-economic use. It is apparent that a training and capacity building in water resource management and flood.

The community identify prohibited areas with regards to sand exploitation from river bed near the point of riverbank rehabilitation as hazard areas. It's therefore necessary that areas of sand exploitation should have the community trained on use of sandbags at riverbank for soil loss prevention. The communities lack adequate information on regulations and use of dykes in river bank management and therefore efforts to carry out. Along the banks of Nzoia River, part of the dyke that had been constructed has been damaged by flood such as floods of 2008. The condition of dykes showed a need of rehabilitation of both southern and northern parts owing to the fact that there're specific weak points on the dyke which are vulnerable to breach should water levels rise beyond the main river channel. For instance, before 2008 floods, flooding frequency due to broken dyke was so low but has gradually increased over time since then. This could be attributed to the fact that many parts of the dyke are weak and therefore need for reinforcement or repair as this has never been the case.

Coping strategies, disaster management training, training on maintenance of dykes, are in great need. Agriculture technology that would avoid flood and drought damage to crops are still necessary as non-structure measures in the management of flood water.

CHAPTER SIX: SUMMARY, CONCLUSIONS AND RECOMMENDATION

6.0 Introduction

This chapter presents a summary of the research results, conclusions and recommendations made towards enhancing conservation and effective management of water catchment areas. The conclusions are drawn from the research findings and discussions. Areas for further research have also been proposed for exploring flooding problems facing water catchments so that they can be addressed appropriately.

6.1 Summary

Rainfall duration and intensity affects flooding by increasing run off since it influences the infiltration and ground seepage rates especially by saturating the soils. Most surges had a tendency to happen in August, which harmonizes with wetter months of the year (July, August, September, April and March) as per the rainfall data.

Morphometric characteristics show that areas with stream orders 1 and 5 have the highest flood risk because their bifurcation ratio is above 2 (Table 2). The areas traversed by stream orders 2, 3, 4 and 6 have a lower bifurcation ratio thus lower chances of flooding. The computation of bifurcation ratio shows that the values ranged from 1.5 and 3.4 which are within Strahler's range. The average value of Bifurcation ratio is 2.2 which is closer to the lower value of 2. Based on bifurcation ratio, River Nzoia basin flows may encounter delayed time to peak and this is a good characteristic for planning and conveying flood forecasts. Analysis of drainage density distribution has shown uniform conveyance with a mean of 0.24-km. This value lacks control values to base judgment. For this limitation the interpretation is limited to spatial variability and frequent occurrence floods is attributed to effect off relief ratio Results on relief ratios and basin orders shows that high relief ratios of above 0.03 are characteristics of upland sub basins and this drops drastically to a low value of about 0.01 for sub basins at the lowlands by suggesting that this reduction in relief indices can increase flow width as flow depth cannot be increased in this zone of the basin due to topographic restrictions

The findings of the study indicate that there has been spatial change in land use over the years in River Nzoia basin. Land use land cover analysis shows that there is increase in size of the bare ground area which is an indication of wetland degradation and a consequent replacement of wetland vegetation with bare soil that is prone to erosion. Increased farming activities along the Nzoia River therefore have resulted to deforestation in search of cultivation areas increasing the vulnerability of the area to flooding. The rapid expansion of urban centres in the lower parts of the catchment (Rwambwa, Kakamega, and Bungoma) can be said to be a major contributing factor to the annual devastating floods.

Various structural and non-structural measures have been put in place, but their efficiency is lacking for various reasons like wear and tear of some dykes. 93.88% agreed to having experienced floods. Although chances of reoccurrence of floods were evident only 40.82% had put in place measures to cope. Only 51.02 % acquired information on floods from the government. Lack of protective dykes (98.33%); lack of repair of the existing dykes (95.92%) and rainfall intensity (79.59%) are the main causes of flooding.

6.2 Conclusion

We can reasonably say that, high rainfall intensity and prolonged duration of rainfall, LULC, and high bifurcation ratio are factors that contribute to flooding. Without proper coping strategies, disaster management training and training on maintenance of dykes results in adverse flood effects.

6.3 Recommendations

Policy makers

The national government should mobilize the community through the chiefs in public barazas and create more awareness on flood management practices.

County government

The county government should link the communal groups and beneficiaries with other stakeholders. In linking people to other societies, it is important to advance partnerships with conjoint benefits where all parties are equal. Memorandums of understandings can be developed to support such partnerships

Researchers

There is need for researchers to undertake periodic studies for land-use land cover changes in flood prone areas of Nzoia catchment and propose appropriate actions

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