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SPECTRAL ANALYSIS OF RAINFALL SERIES IN KENYA

By

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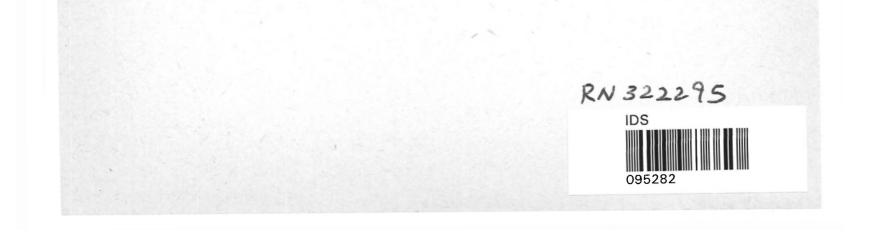
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SPECTRAL ANALYSIS OF RAINFALL SERIES IN KENYA

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Tichaendepi R. Masaya

ABSTRACT

Rainfall series in various parts of Kenya is subjected to spectral analysis. Contrary to past investigations, it is found that no rainfall behaviour at any rainfall station possesses a ten-year cycle neither is there any important five-year cycle. Most rainfall series in central Kenya possess an average cycle of between $2\frac{1}{2}$ to $3\frac{1}{2}$ years while those on the Coast and Lake Victoria area have a 2.2-year cycle. There are secondary cycles of various lengths for most stations.

SPECTRAL ANALYSIS OF RAINFALL SERIES IN KENYA

I. Introduction

Public policy makers and farmers in Kenya are concerned with rainfall variation. Long droughts have been responsible for crop failure and deaths of both domestic animals and wild life. On the other hand heavy rains following these droughts have caused flood damage in certain areas.

A study of rainfall variation was done by Lumb (1966) and Morth (1967) using rainfall data of the Lake Victoria catchment area. The data was analyzed in the time domain using autocorrelation functions of the rainfall series. They both discovered a 10 year cycle. This suggests that rainfall in most parts of Kenya would tend to possess this cycle due to common factors influencing rainfall behaviour in this country.

Obasiand Rodhe (1974) subjected annual rainfall series (averaged over 9 stations in central Kenya) to spectral analysis. Two prominent peaks of equal power were found. The two peaks had average periods of 3 and 5 years. Strictly speaking spectral power corresponding to the latter was stronger. There was a third peak with a period of about 20 years. This peak was, however, not very prominent. Failure to use spectral windows such as those suggested by Tukey, Parzen and others, was responsible for the artifical peaks in the spectrum.

The object of this paper is to investigate the validity of the existence of these periodicities using spectral methods of analysis.

2. A Review of the Theory of Spectral Analysis.

Time series data are formed by recording results of an

experiment at equidistant points in time in an ordered form,

say x_1, x_2, \ldots, x_T and the set of the s

The time series will exhibit widely fluctuating properties and the series itself is a function of time.

In studying the behaviour of a time series, statisticians usually decompose the series into a deterministic component (the trend), a cyclical component, a seasonal component and a purely random component. To gain detailed knowledge of the phenomena described by the series it pays to study each of these movements in isolation.

The trend and the seasonal component are deterministic and once they are known they are of little interest as long as conditions remain sufficiently stable.

The trend distorts spectral estimates near the lower frequency area. For monthly data, excessive power at seasonal frequenncies may obscure the existence of peaks at nonseasonal frequencies. Further, the existence of a trend may prevent the autocorrelation function from falling to zero at appropriate lags when the series is studied in the time domain. Moreover, spectral methods are only applicable to stationary series.

The first step, therefore, is to detrend and deseasonalize the data. For a single series the autocovariance function is given by

(2.1) $C_{XX}(\tau) = \frac{1}{n} \frac{n \overline{\Sigma} \tau}{t-1} (X_t - \overline{X}) (X_{t+\tau} - \overline{X}),$ where $\overline{X} = \frac{1}{n} \frac{n}{\Sigma} X_t$ is the mean of the time series,

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τ = lag or time delay,

n = the number of the observations under study.

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It is sometimes useful to reduce the covariance function to its dimensionless form. This is achieved by dividing (2.1) by the variance to obtain the following autocorrelation funcation: $(2, 2) = P_{-1}(x) - C_{-1}(x)$

, and the constraint of $\left\lfloor \frac{1}{3} \right\rfloor$ - in this ways that is

(2.2)
$$R_{xx}(\tau) = C_{xx}(\tau)$$

 $C_{xx(0)}$

A plot of $R_{\chi\chi}(\tau)$ against τ is called a correlogram.

Alternatively, a time series may be described in the frequency domain by the Fourier transform of the autocorrelation function given by

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$$(2.3) \quad f_{XX}(\omega_{j}) = \frac{1}{2\pi} \frac{M}{\tau^{2} - M} R_{XX}(\tau)\lambda(\tau) e^{-\omega_{j}\tau}$$

$$= \frac{1}{2\pi} \left[R_{XX}(0) \lambda(0) + \sum_{\tau=1}^{M} R_{XX}(\tau)\lambda(\tau) \left\{ e^{i\omega_{j}\tau} + e^{-i\omega_{\tau}\tau} \right\} \right]$$

$$= \frac{1}{2\pi} \left[1 + \sum_{\tau=1}^{M} R_{XX}(\tau)\lambda(\tau) \cdot \left\{ \cos\omega_{j}\tau'' + i\sin\omega_{j}\tau + \cos\omega_{j}\tau - i\sin\omega_{j}\tau \right\} \right]$$

$$= \frac{1}{2\pi} \left[1 + 2 \frac{M}{\sum_{\tau=1}^{\Sigma} R_{XX}(\tau)\lambda(\tau) \cos\omega_{j}\tau} \right]$$

$$= \frac{1}{2\pi} \cdot \frac{1}{\pi} \sum_{\tau=1}^{M} R_{XX}(\tau)\lambda(\tau) \cos\omega_{j}\tau$$

where $f_{xx}(\omega_j)$ = estimate of normalized power spectrum averaged

over a frequency band centered at ω_{i} ,

 $(-1)^{-1}$

- M = the number of frequency bands to be estimated
 - or the truncation point.
- $\lambda(\tau)$ = is a lag window = 1 for $\lambda(0)$.

 $2\pi j$ πj

(2.4.)
$$f_{xx}(\omega_j) = \frac{R_{xx}(0)\lambda(0)}{2\pi} = \frac{1}{2\pi}$$
 is a flat white noise spectrum.

Apart from suggesting a model for the available data, spectral methods have certain advantages over correlation methods. These advantages are discussed in some detail by authors such as Granger and Hatanaka (1964), Nerlove (1964), Jenkins and Watts (1968) and many others.

If monthly data are used their spectrum will exhibit narrow and sharp peaks at seasonal frequencies if the data possess a seasonal movement. Seasonal frequencies, periods, and the number of times a cycle is completed in a year are given in Table 1.

49 201202

Table 1 Seasonal Period Number of Frequency (Months) times cycle (cycles per month) completed in a year 0.083 12 1 0.167 2 6 0.250 3 0.333 3 . 4 0.417 5 2.4 0.500 2 6

3. The Data

Monthly rainfall data recorded at Kabete station from 1916-1973 was used in the analysis. The data are given in Table 2. Annual data for the same station was available from 1910-1973. These are shown in Table 3 together with

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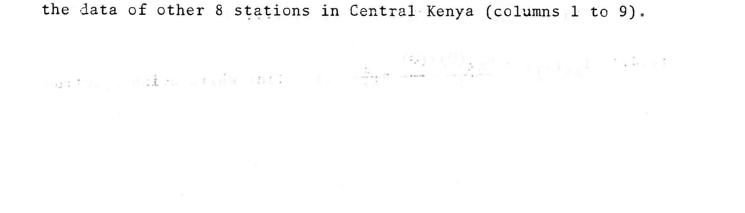


TABLE 2 KABETE MONTHLY RAINFALL (millimetres)

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			J	r	м	A	М	Month J	J	A	S	0	N	D
		lear												
		L916	83	55	118	317	170	33	0	41	97	35	152	70
			163	90	223	323	200	154	18	68	53	120	40	3
		L918	2	8	11	242	103	29	27	2	17	85	59	67
		L919	3	114	70	175	129	31	62	40	0	81	115	21
		L92 o	60	0	248	332	107	34	6	11	41	86	183	171
		1921	24	48	37	224	39	9	9	19	0	27	130	85
		L922	97	86	265	342	119	16	17	37	56	97	58	152
		L923	55	110	242	374	419	59	24	8	15	53	165	34
]	L92 4	0	27	138	171	8 o	22	7	50	3	18	139	57
		L925	181	44	136	84	73	33	20	28	ο	33	114	66
	10 N. I	L926	38	29	14 4	231	152	26	13	34	67	66	158	15
		L927	27	24	213	195	129	43	14	ο	o	53	73	59
		L928	19	6	45	125	350	47	l	12	5	68	169	46
		L929	15	l	3	371	148	3	49	29	71	77	137	189
		L930	74	82	262	387	334	23	12	4	8	70	155	37
		1931	28	52	173	394	122	44	16	20	lo7	102	104	108
		.93 2	74	36	189	299	169	27	24	. 16	19	52	92	158
	· · ·		144	0	37	56	58	7	17	54	lo	52 65	92 123	85
		1934	3	25	33	126	193	49	31					1
										10	0	65 5 3	48	67
		L935	0	150	48	108	121	62	0	53	12	53	144	117
		L936	96	142	148	218	53	67	6	12	14	85	89	104
3.1.		1937	8	1	191	331	331		4	.7	116	146	243	96
		L938	30	56	125	105	lol	7	9	9	38	37	149	118
		939	22	17	138	233	46	9	13	17	. 17	27	59	6
	د ₁	940	49	66	106	271	300	38	41	14	l	50	92	37
	1	L941	35	113	144	343	187	2 85	2	12	l	7	89	190
		L942	21	42	245	269	221	87	28	36	l	30	41	114
	i	L943	1	80	44	184	65	44	2	20	lo	9	77	57
		L944	2 2	0	90	172	33	21	23	8	64	65	131	70
		L9 4 5	43	74	57	43		lo4	69	37	24	ı	155	29
		1946	lo	3	21	147	145	15	11	70	138	217	66	48
			lol	28	180	343	332		32	2	48	8	79	57
		1948	9	6	98	210	122	77	3	39	23	69	93	200
		194 9	ì	38	17	260	58	lo	3	24	18	10	63	123
		L950	38	19	114	254	73	22	8	45	9	74	81	11
		L951	11	4	100	352	315		32	24	3	103	4 4	225
				24	29	376	207		11	6	44	22	- + 4 79	24
		L952	16											1
	4	L953	33	0	45	135	198	42	17	34	36	53	132	34
	1	L954	37	76	1	303	297	26	36	、17 17	5	34	113	28
		L955	54	75	33	207	133	3	17	15	28	66	58	220
0	b = = =	L956	202	91	99	175	136	9 -		10	10	22	32	42
		1957	157	. 25	43	201	450		11	6	51	65	175	87
~		1958	25	160	36	141	346		93	1	8	lo	55	55
		1959	20::	79	116	78	130		15	54	21	41	288	57
		1960	41	7	248	192	109	33	7	lo	32	79	76	33
		1961	3	13	76	175	152	28	33	38	48	178	549	371
		1962	201	43	41	180	302	58	3	41	io	86	117	127
	:	196 3	87	55	103	385	465	29	5	65	19	7	176	319
		1964	50	120	56	436	135	17	28	70	12	29	78	94
	. :	1965	75	8	23	322	80		14	lo	7	°895	143	163
		1966	147	51	203	220	112		4	64	19	46	165	48
		1967	0	9	24	285	488		26	44	44	138	136	1
		1968	0	123	210	246	185		6	44 6	44 6	130 52		40
		1969	82	90	104		119						285	163
		1989 1970				40			6	40	9	33	111	14
			132	3	132	423	166		lo	6	10	29	131	35
		1971	68	11	27	235	300		27	20	20	24	74	134
	-	1972	17	83	57	28	17 1		17	8	49	198	146	31
		1973	140	65	8	232	43	42	4	9	62	19	107	41

Source: East African Meteorological Department

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				ANNUAL	TABLE ; RAINFALL(m:		a):				
(1) (2)	(3)	(4)	<u>CENTRA</u>	(6)	KISUMU, 1	LODWAR AN.	MOMBAS	<u>.</u> (10)	(11)	(12)	(13)
(1) (2) EAR EMBU 891 892 893	KABETE	KIAMBU	MACHÁKOS	MERU	nuranga	NARURU	NGÒŃĠ	NYERÍ		. KISUMU	LÒDWÁR
1895 1995 11109 111109 111109 111111 111111 111111 111111 111111 111111 111111 111111 111111 111111 111111 111111 111111 111111 111111 1111111 <tr< td=""><td>7149150152191 28223033565 8856435853961695 999492531</td><td>9967118944719877 418596668897600511632199056795503168954793964</td><td>$\begin{array}{c} 1066\\ 1316\\ 6595\\ 14804\\ 10085\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1296\\ 1315\\ 1296\\ 1315\\ 1296\\ 1315\\ 1297\\ 1296\\ 1315\\ 1297\\ 1296\\ 1257\\ 1297\\ 1296\\ 1257\\ 1297\\ 1297\\ 1296\\ 1257\\ 1297\\ 1297\\ 1297\\ 1296\\ 1257\\ 129$</td><td>$\begin{array}{llllllllllllllllllllllllllllllllllll$</td><td>$\begin{array}{c} 1123123997176902333677\\ 11369387623346562346556694831330927019267097508858\\ 11212239333997176902333677\\ 113693876223346556234655694831132009750997508857\\ 112122393297508877\\ 11368528762346556994831132009750997508858\\ 1142873399388762234655699483313399382769217993877508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989835\\ 11428733993876996217989750997508835\\ 11428733993876999750997508835\\ 11428733993876996217989750997508835\\ 11428733993876999750997508835\\ 114287339938769997509975098835\\ 114287339938769997509975098835\\ 114287339938562334655699483313993885\\ 1142873399387699975509975509975508835\\ 114287339938576999755099755098835\\ 11428733993857699755099755098835\\ 11428733993857699755099755098835\\ 11428733993857699755099755098835\\ 1142873993855669966696669666666666666666666666$</td><td>6293639911762399111782996536796986948934005593772986980755869980755869980755869980755869980755937729678698071578011758011</td><td>572776913916 92869676790244523621411642702835539161454525 8885686377383888 60458370244523622141164270283553916145454525</td><td>612255858445369205512 81970897408974999641204737037035093978011786880725 10759730689205512 108919976917370370350935093978017789370035093978011786680725</td><td>79985694682474 371111117116 13871898076922220145 8029989298399925640728 1111187901647692220145 802998399298399925640728 879874145 802998399298399925640728</td><td>10350370146611095682 50487702695067889490676821021325400735037949395687</td><td>4497792513833395288966552106339473733 19831882 19894773733 154330 155288 259477373 154330 1689 16956 16956 1498 16956 1498</td></tr<>	7149150152191 28223033565 8856435853961695 999492531	9967118944719877 418596668897600511632199056795503168954793964	$\begin{array}{c} 1066\\ 1316\\ 6595\\ 14804\\ 10085\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1285\\ 1296\\ 1315\\ 1296\\ 1315\\ 1296\\ 1315\\ 1297\\ 1296\\ 1315\\ 1297\\ 1296\\ 1257\\ 1297\\ 1296\\ 1257\\ 1297\\ 1297\\ 1296\\ 1257\\ 1297\\ 1297\\ 1297\\ 1296\\ 1257\\ 129$	$\begin{array}{llllllllllllllllllllllllllllllllllll$	$\begin{array}{c} 1123123997176902333677\\ 11369387623346562346556694831330927019267097508858\\ 11212239333997176902333677\\ 113693876223346556234655694831132009750997508857\\ 112122393297508877\\ 11368528762346556994831132009750997508858\\ 1142873399388762234655699483313399382769217993877508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989750997508835\\ 11428733993876996217989835\\ 11428733993876996217989750997508835\\ 11428733993876999750997508835\\ 11428733993876996217989750997508835\\ 11428733993876999750997508835\\ 114287339938769997509975098835\\ 114287339938769997509975098835\\ 114287339938562334655699483313993885\\ 1142873399387699975509975509975508835\\ 114287339938576999755099755098835\\ 11428733993857699755099755098835\\ 11428733993857699755099755098835\\ 11428733993857699755099755098835\\ 1142873993855669966696669666666666666666666666$	6293639911762399111782996536796986948934005593772986980755869980755869980755869980755869980755937729678698071578011758011	572776913916 92869676790244523621411642702835539161454525 8885686377383888 60458370244523622141164270283553916145454525	612255858445369205512 81970897408974999641204737037035093978011786880725 10759730689205512 108919976917370370350935093978017789370035093978011786680725	79985694682474 371111117116 13871898076922220145 8029989298399925640728 1111187901647692220145 802998399298399925640728 879874145 802998399298399925640728	10350370146611095682 50487702695067889490676821021 3 25400735037949395687	4497792513833395288966552106339473733 19831882 19894773733 154330 155288 259477373 154330 1689 16956 16956 1498 16956 1498

1973 638 772 620 804 1000 720 727 572 962 757 212 120 1974

Source: East African Meteorological Department.

An areal average for the 9 stations in Central Kenya is shown in column (10) of the table. The longest series in Central Kenya is that of Machakos and Muranga whose recording started in 1894 and 1901 respectively.

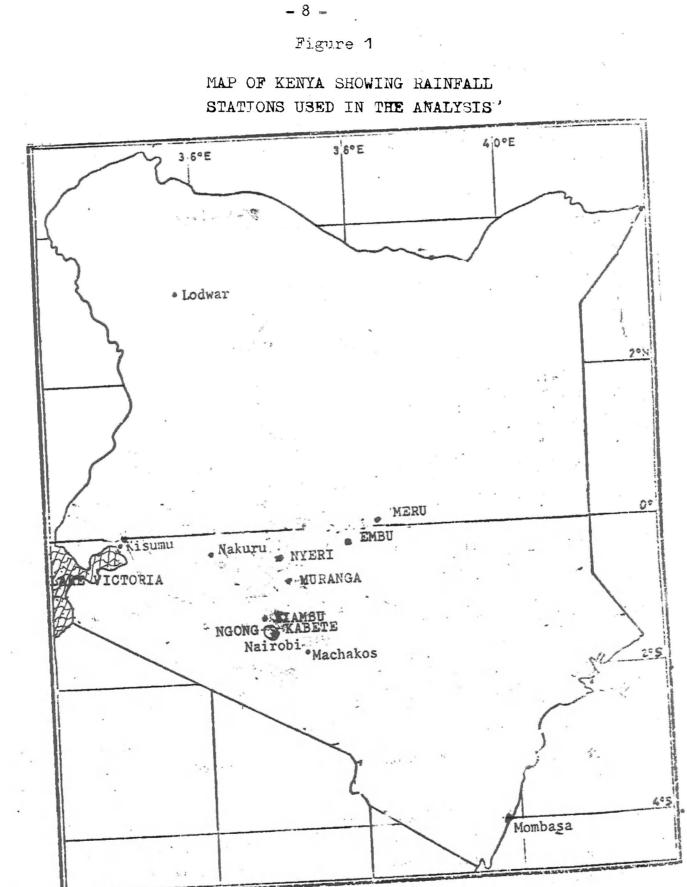
It was also necessary to study annual rainfall variation on the East Coast, the arid North and the Lake Victoria area. Kisumu station (column 11) was chosen to represent the Lake Victoria region. This station has fairly long data starting from 1903 to 1974. The data for 1955 was not recorded and the value used for that year was an estimate. In addition the annual data for the Lake Victoria catchement was also used. This data was used by Morth (1967) for the period 1899-1937. The data was estimated by regression methods over a very long period and is likely to be in errors. Further, although this series was used it is rather too short for spectral analysis. It is not shown here.

The arid North is represented by annual rainfall data read at Lodwar from 1922 to 1974 (column 12). There was no data for the war years 1939, 1940 and 1941. The values for these years were estimated.

Mombasa representing the coastal area has the longest and uninterrupted data of all the stations. The data for this station starts from 1891 to 1973. These are given in column 13 in Table 3.

The geographical position of the meteorological stations mentioned above are shown in Figure 1.

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. . . 4. Spectral Analysis of the Data

The spectral window used was Parzen's with a trunction point M = 50 for monthly series and M = 15 for all annual series.

The first data to be subjected to spectral analysis was the Kabete monthly series in their original form. The spectrum in Figure 2 shows very high power in the low frequency area represented by frequencies less than 0.03 cycles per month or periods greater than 33 months. This implies the existence of an upward trend in the data over the sample period. The power densities at low frequencies is so domineering that peaks at seasonal frequencies, corresponding to those in Table 1, have been considerably reduced or disappeared altogether.

To remove excessive power at low frequency the data were prewhitened by the method of first differences. The spectrum of the pre-whitened series is in Figure 3. Not only are peaks prominent at the majority of seasonal frequencies, they have also become narrower and sharper as expected. The peak with an average period of 6 months, contributing about 68% to total variance, is the most important. The 12 month cycle contributes less than 14% and is the least important of the existing peaks. There are no peaks at the frequency band centred at 0.417 cycles per month.

When the Kabete annual series was subjected to spectral analysis power density was so dominant at low frequencies that no peak was found in the middle frequency band nor in the high frequency band. As a result it was decided to pre-whiten all the data used in the rest of this paper.

1 See Appendix

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Figure 4 shows the normalized spectral density function for the pre-whitened Kabete annual series. The spectrum is dominated by a peak associated with a 3.2 year cycle. There is a less important 5-year cycle corresponding to the frequency band centred at 0.2 cycles per year or 2 cycles per decade.

The spectrum of the areal average of nine stations in Central Kenya is different from the Kabete spectrum in that it possesses an extra peak associated with a 2.2 year cycle as shown in Figure 5. This cycle is more important than the 5-year cycle. However, the most important cycle in the spectrum is the one corresponding to the 3.2 year cycle.

The dissimilarity between the spectrum of the average series and that of Kabete which is a part of the areal average demands that a few of the stations in the area be studied individually. Additional stations studied, were Machakos, Meru and Nakuru.

The spectrum of Machakos is in Figure 6. There is a "strong" peak corresponding to a frequency band centred at 0.2 cycles per year. This indicates a dominant cyclical component whose period is 3 1/2 years. There is a subsidiary peak cprresponding to a cycle of 2 1/2 years. Spectral power drops rapidly at frequencies lower than 0.2 cycles per year and those higher than 0.4 cycles per year.

Like the Machakos spectrum, the Meru spectrum in Figure 7 does not possess a 5-year cycle. Rather, the spectrum has one peak lasting for an average period of 2.6 years. At high frequencies the spectrum is flat. This indicates the existence of White noise at frequencies higher than 0.45 cycles per year.

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The last of the stations to be analysed in central Kenya is Nakuru whose spectrum is in Figure 8. The 2.5 year cycle contributes about 31% to total variation, while the 3.1 year cycle contributes 29%. There are no peaks at frequencies lower than 0.325 cycles per year.

Now it seems clear that the 2.2 period in the spectrum of the areal average is a result of the Machakos' 2.6, Meru's 2.5 and Nakuru's 2.5 year cycles, together with similar cycles of the stations not analysed individually. The 3.2 year cycle is a combination of Machakos' 3.5, Kabete's 3.2 and Nakuru's 3.2 year cycles. The 5 year cycles is a reflection of the 5 year cycles in the Kabete spectrum.

The spectrum of Lake Victoria catchment area is shown in Figure 9. There is an important peak indicating a cylical component with a period of 3.2 years. A secondary peak exists corresponding to a cycle of 2.2 years. As mentioned earlier these data cannot be relied upon. To check on the realities of the periodicities of these data, the data recorded at Kisumu PC's office was analysed. Rainfall at this station should be influenced by the factors influencing rainfall in the Lake Victoria area in general. The Kisumu spectrum in Figure 10 is similar to that of the Lake Victoria catchment area in that both spectra have a peak corresponding to a 2.2 year cycle and dissimilar in that there is no 3.2 year cycle in the Kisumu spectrum. There is a negligible contribution at the peak with a 5 year period.

Mombasa spectrum (Fig. 11) like Kisumu has a dominant 2.2

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year cycle with a peak contributing to 33% of total variance. There is also a second peak with a period of 3.2 years. Figure 12 shows that the most important peak for Lodwar rainfall series is the one corresponding to a 2.3 year cycle. A 3.3 year cycle is also important. There is no sign of a period longer than this.

Table 4 is a summary of the results of the above analysis. The table also provides a contrast of the spectra by the number of peaks and by distinguishing important cycles by the amount of contribution to total variance by components of various periods. The dominant peaks for Kisumu, Mombasa and Lodwar correspond to cycles of about the same periods.

Table 4

Contribution to Variance by Components Corresponding to Peaks

Period (Years)		Areal Ave.	Machakos	Meru	Nakuru	Lake Victoria		Mombasa	Lodwar
2.2		26%				29%	32%*	1 3.38*	ala e
2.3		14,022			. A. H.	i dan di	$p^{i}=a^{ij}$		32%*
2.5	S. S		23%	n, 4° ,	31%*	248.29		-00 - 11 - 1 	
2.6		a m		28%*					Salara
3.1				- y- 18	29%	x 148 - 1	ст. <u>.</u>	d e se	e. El 2 y c
3.2	30%*	28%	* 55 200	1 ^{- 1} 12	i in	35%*	ant 1	23%	ા ગે નામા
3.3		ten da	1 - E	15	ar grit	at que	1		29%
3.5	1997 - L	a de des	28%*		v : c	an an shiri Shiriba			
5.0	19%	20%					12	() j.	

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* most dominant peak

- A.A.

- 12 -

Usually stability is achieved by widening the frequency band although this is done at the expense of fidelity or resolution. Band width is inversely related to the truncation point. Thus the reduction of the truncation point may lead to the disappearance of the weak peaks in the spectrum.

Figure 13 shows the nine spectra resulting from the reduction of M from 15 to 6. Most spectral power is now concentrated at high frequencies. All spectra are fairly smooth below 2.9 cycles per decade.

Table 5 is a summary of the results of these spectra. Each spectrum has one peak corresponding to periods in the first comlumn of the table. A station possessing that period is indicated by a star.

The Kisumu and Mombasa spectra are very similar. Both have a 2 year cycle. In the other hand Meru, Nakuru and Lodwar are now alike in that they all possess a cycle of a period ranging from 2 to 2.2 years. Lake Victoria catchment area has a cycle exceeding that of Kisumu by almost two months. The three-year cycles for Kabete and the nine stations in Central Kenya, as well as the 3 1/2 year cycle for Machakos have peaks showing up very clearly even at this low truncation point.

Table 5

Kisumu Mombasa Lodwar Period Kabete A real Machakos Meru Nakuru Lake Victoria (Years) Ave. * * 2 * * 2-2.2 × * 2-2.4

Important peaks when M = 6

- 13 -

2.7-3	*	*		1		 	
3.4			*				
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5 Conclusion

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The spectra of all the rainfall series reveal that high frequencies contribute much more to the variance of the process than do the low frequencies. There is, therefore, no possibility of cycles of more than 5 years.

In central Kenya Meru, and Nakuru rainfall patterns possess an approximately 2 1/2 year cycle. The latter, however, has a secondary peak corresponding to a cycle of about 3 years. Machakos has two cycles, one dominant cycle of 3 1/2 years and a less dominant cycle of 2 1/2 years. Kabete has a definite 3.2 year cycle. A second peak lasting 5 years is insignificant.

Kisumu, Mombasa and Lodwar have dominant 2.3 year cycles approximately. The last two of these rainfall stations have secondary peaks lasting for a period of approximately 3.3 years. The similarities of the 3 stations may be ascribed to their proximity to large masses of sea and lake waters. The spectrum for the Lake Victoria catchment area is dissimilar from the spectra of rainfall stations near large masses of water in that the dominant peak occurs with a period of 3.2 years as opposed to 2.3 years. Its secondary peak rather than correspond to a 3.2 year cycle, corresponds to a 2.3 year cycle. This is the reverse of the other stations in similar situations. This may be ascribed to three causes.

 That the series is too short to have any meaningful application of spectral methods.

2. The data was estimated by regression methods for too long

a period creating uncertainties about the accuracy of the

estimates.

3. The averaged data may have included a station (or stations) with a different rainfall pattern.

What is interesting however, is that the 10 year cycle claimed by earlier research workers, using these data, does not exist. The maximum period found is about 3 years. Even this is doubtful when compared to the dominance of the 2.3 year cycle found in the spectra of the other 3 stations in similar conditions.

This investigation has further revealed that if the study of periodicities in the rainfall series is carried out for the purpose of policy-making, the use of areal averages may be misleading. Rather, individual analysis of rainfall behaviour at a particular station is more appropriate.

Rainfall behaviour in Kenya suggests, therefore, that irrigation schemes capable of combating drought for periods upto about 4 years are required. Areas near masses of water require shorter periods. However, for the more arid areas of the country such as Lodwar irrigation schemes lasting for infinite periods are required since the highest amount of rainfall recorded here is insufficient for most useful crops. The use of land surrounding lakes for agricultural purposes must be encouraged not only because of irrigation facilities the lakes provide but also because of the shorter periods of drought that exist near these surroundings.

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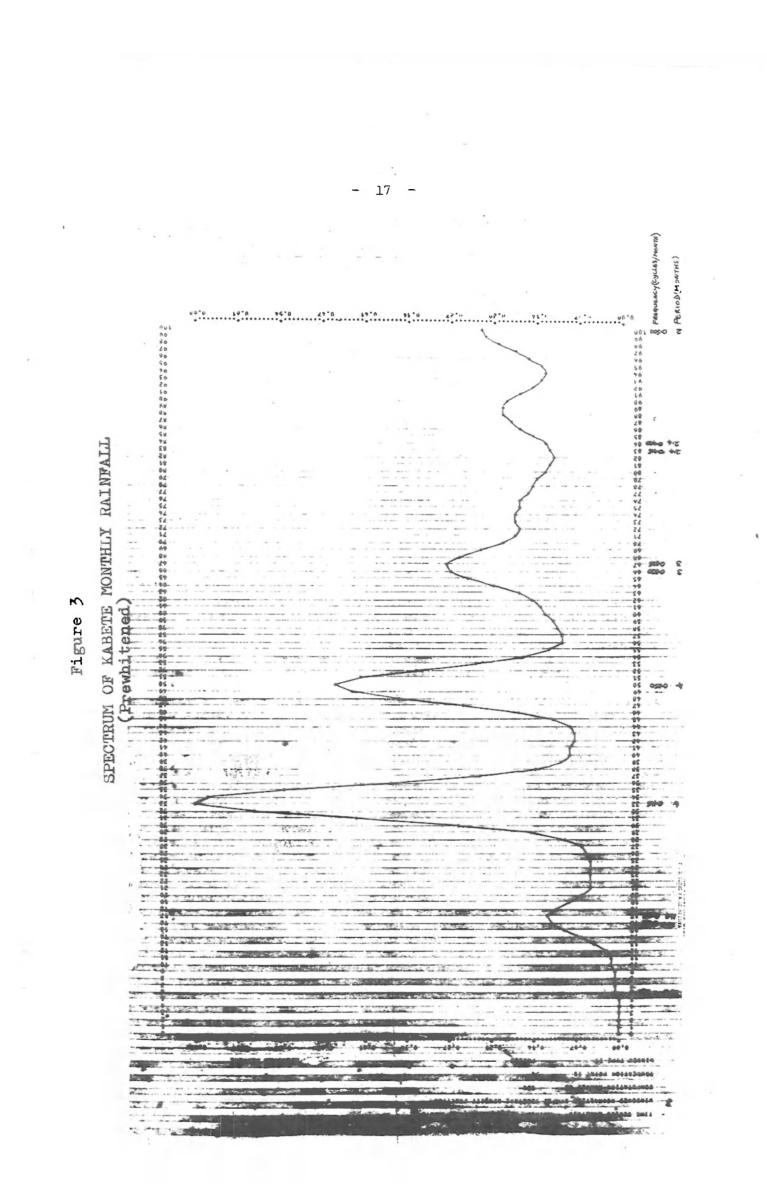
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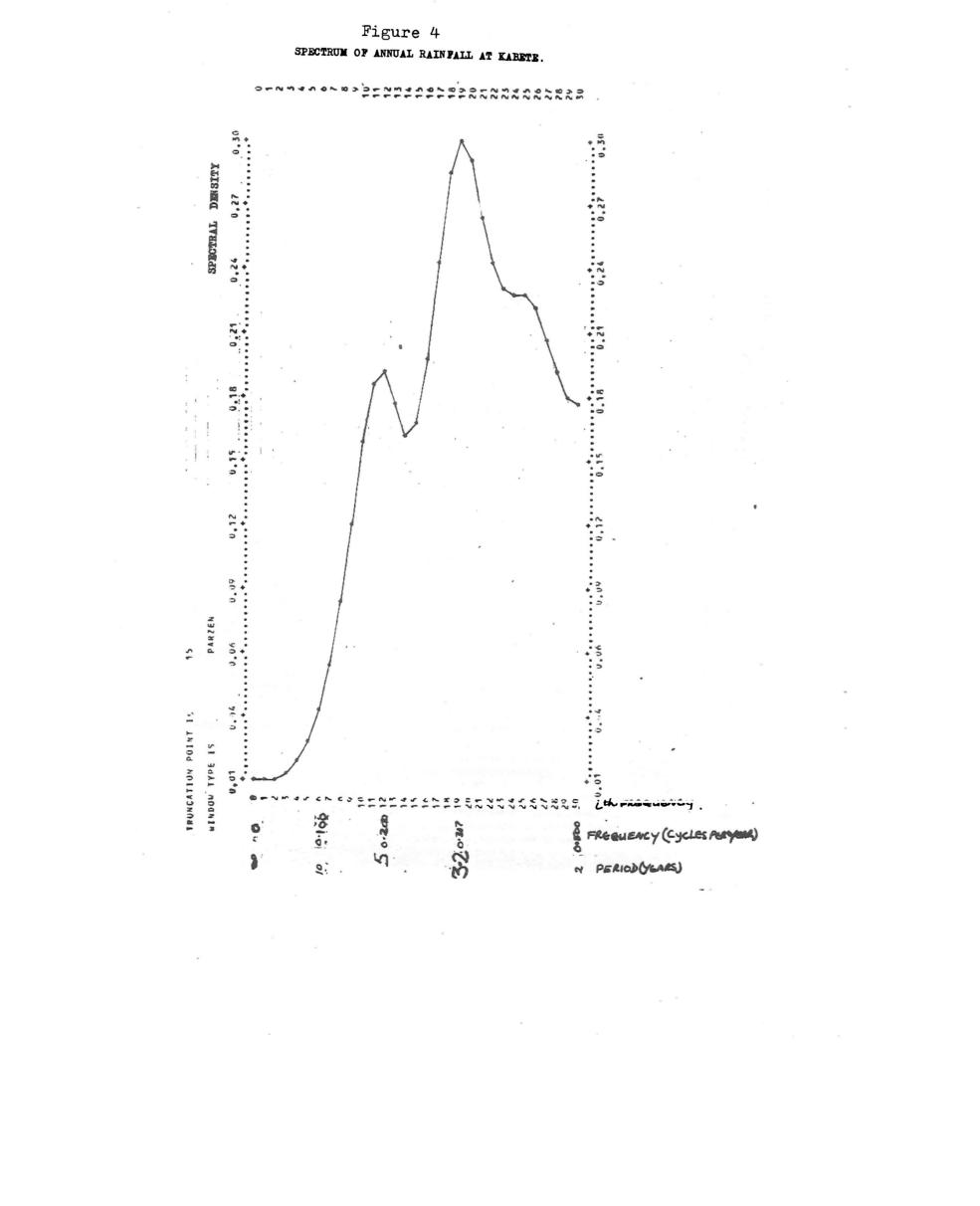
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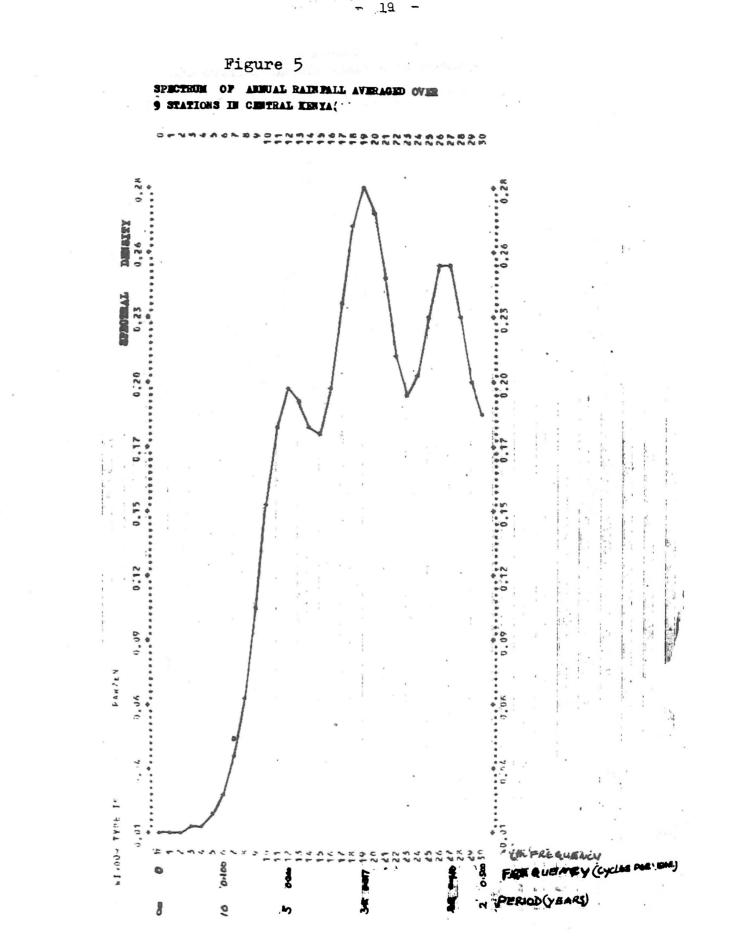
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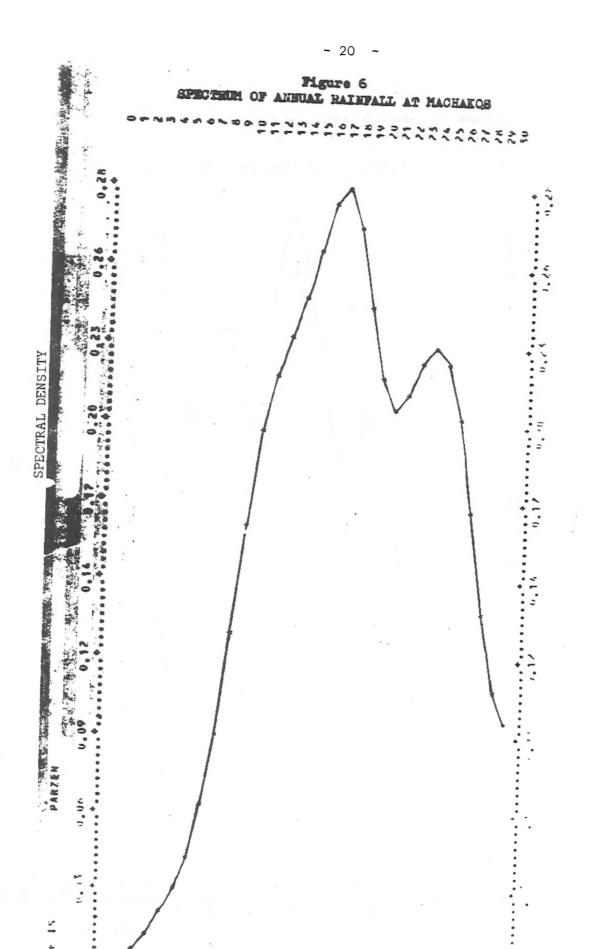


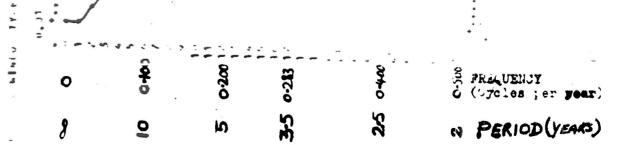
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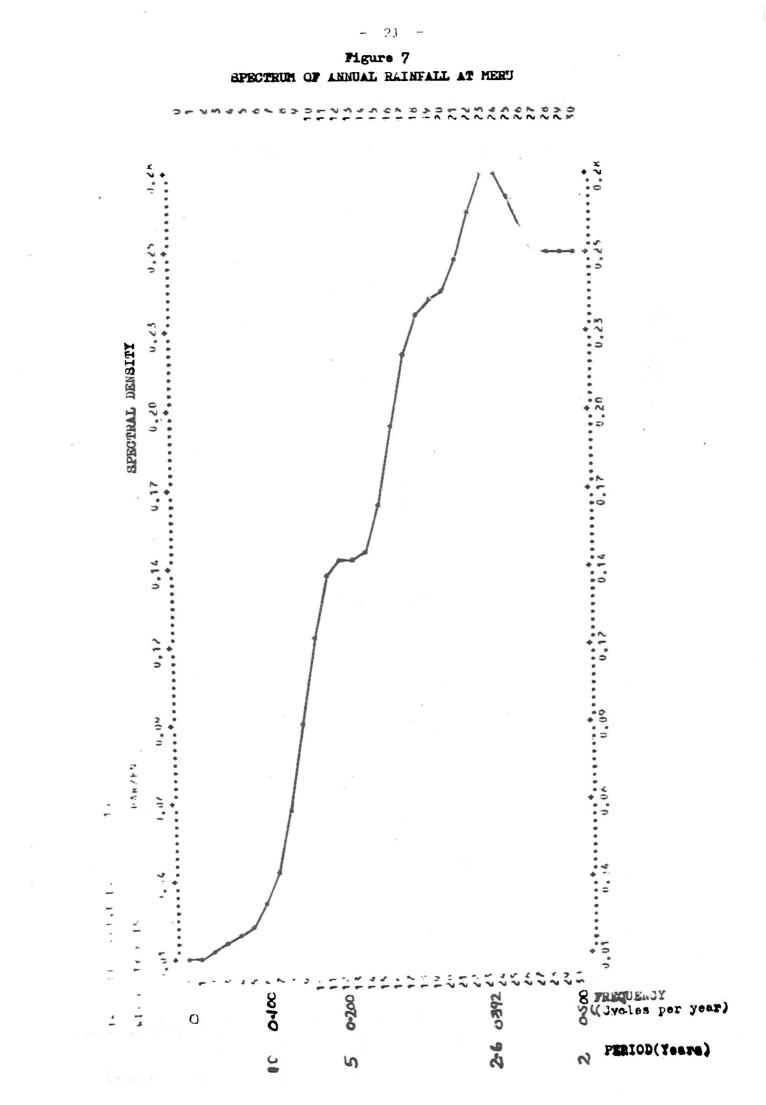


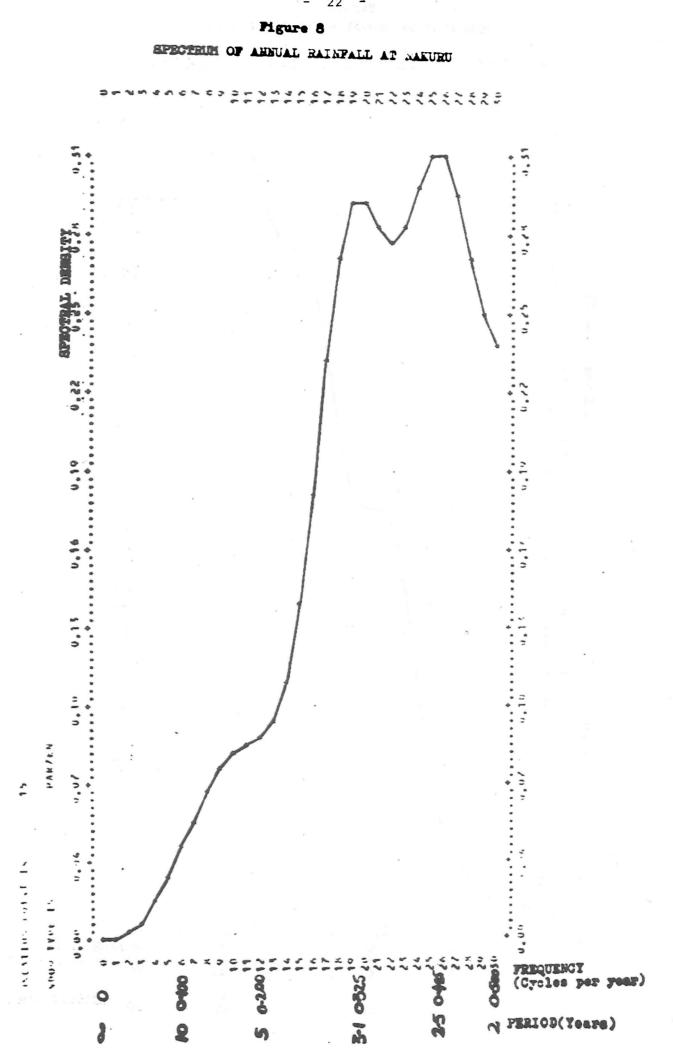
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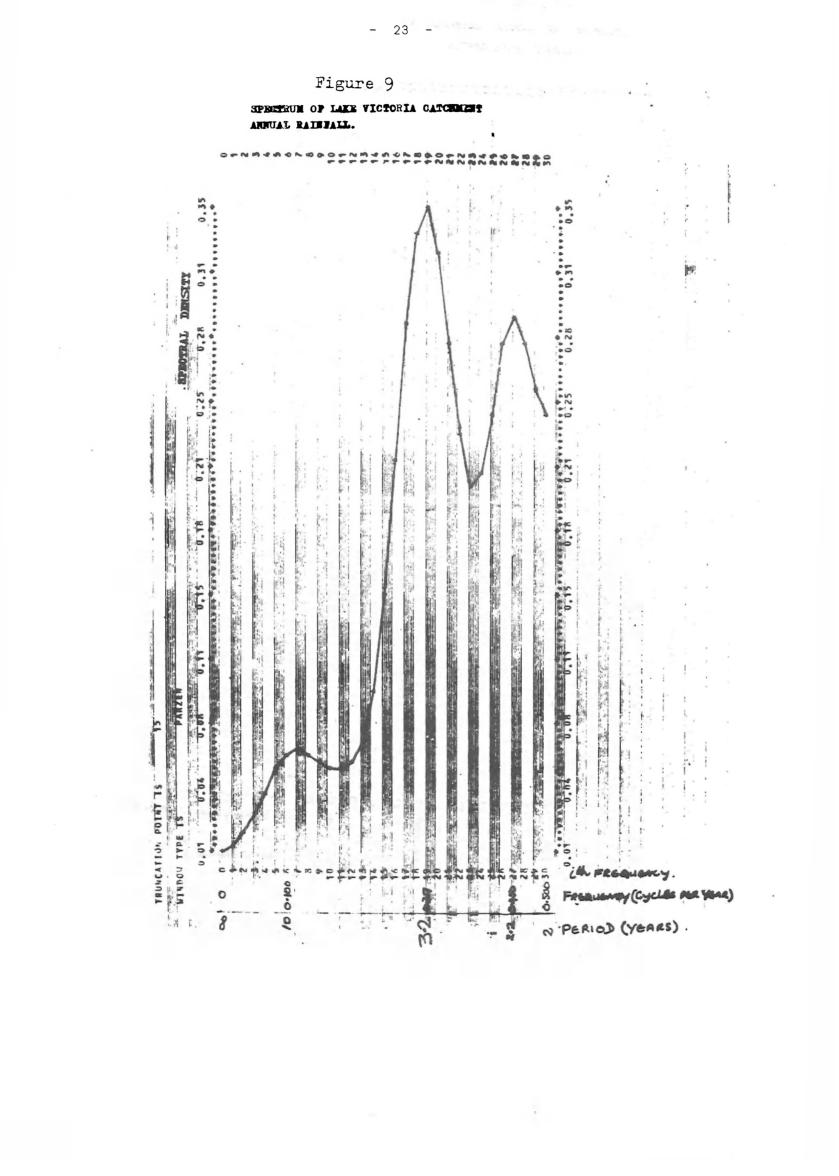


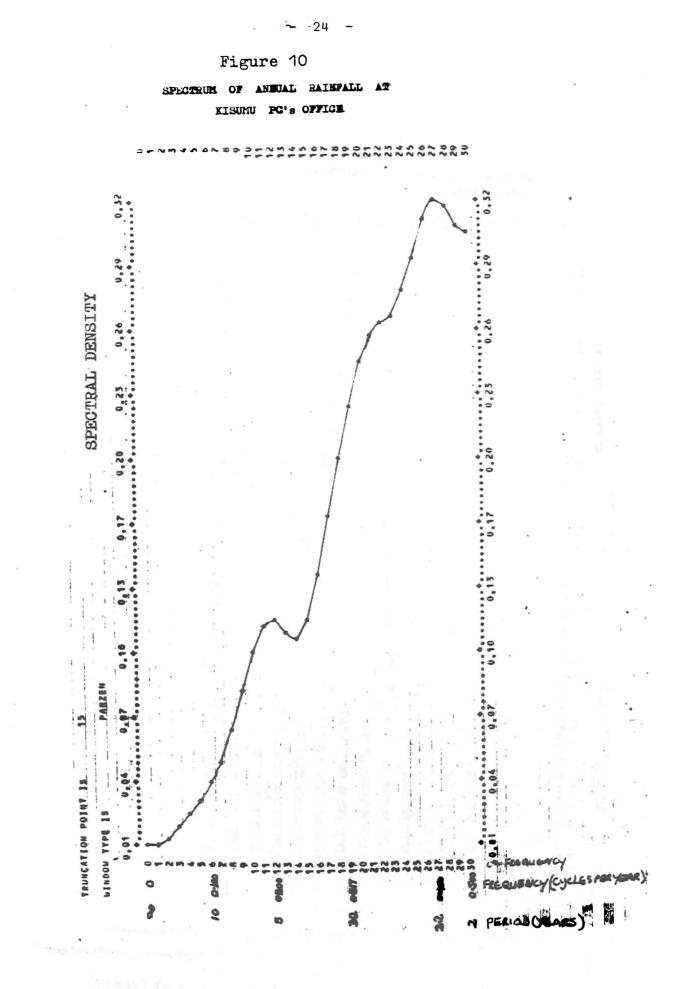


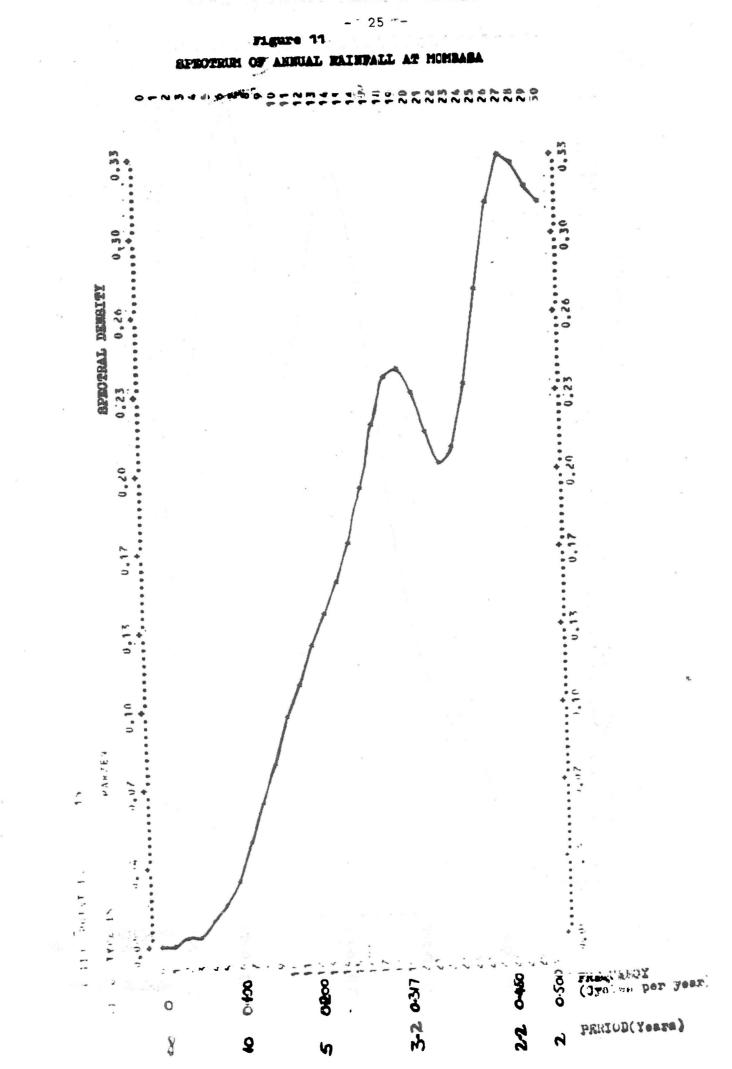


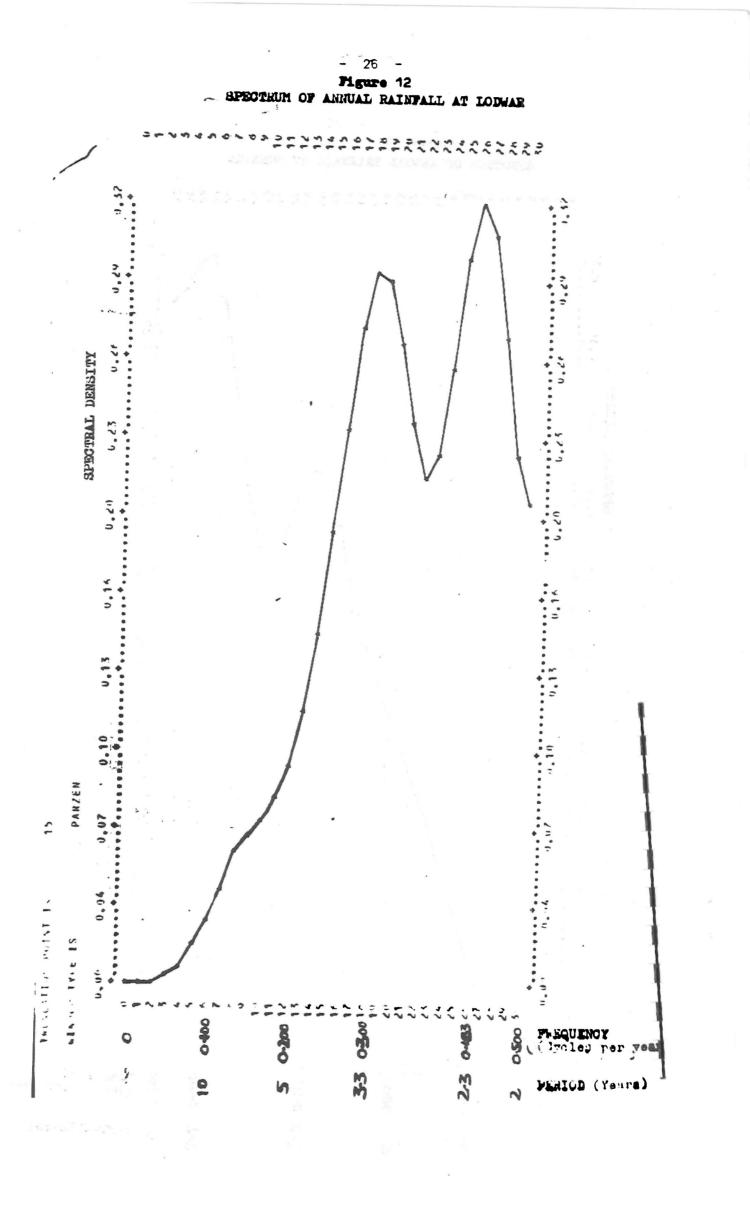
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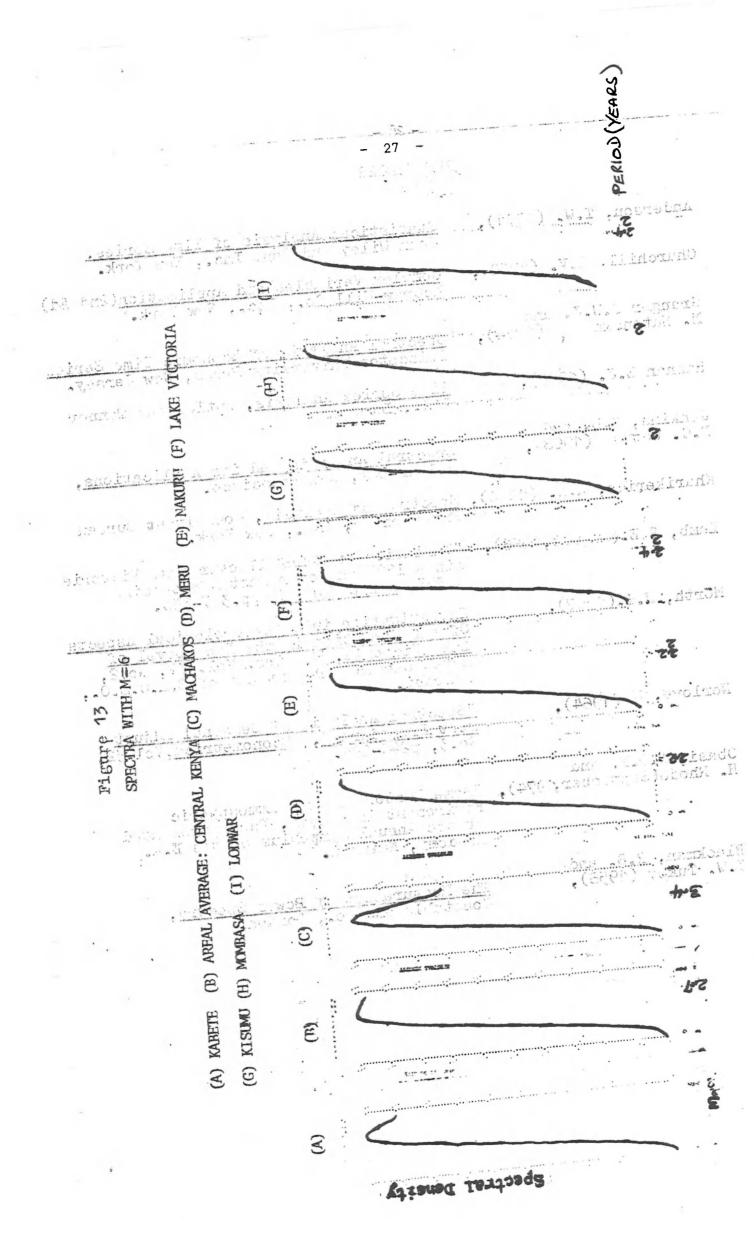
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REFERENCES

Anderson, T.W. (1971),	Statistical Analysis of Time Series. John Wiley and Son, Inc., New York.							
Churchill, R.V. (1960),	Complex Variables and Application(2nd Ed) McGraw-Hill Co., Inc., New York.							
Granger C.W.J. and M. Hatanaka (1964),	Spectral Analysis of Economic Time Series Princeton University Press, New Jersey.							
Hannan E.J. (1960),	Time Series Analysis, Butler and Tanner Ltd., London.							
Jenkins, G.M. and D.G. Watts (1968),	Spectral Analysis and its Applications, Holden-Day, San Francisco.							
Kharikerich, A.A. (1960),	Spectra and Analysis, Consultant Bureau Enterprises, Inc., New York.							
Lumb, F.E.(August,1966),	"Variation of Rainfall over Lake Victoria since 1899 and over East Africa since 1937", <u>Kenva Coffee</u> , pp.347-350.							
Mörth, H.T.(1967),	Investigation into Meteorological Aspects Of the Variation in the Lake Level of Lake Victoria, Memoirs, Vol. IV, No.2, Meteorological Department of E.A.C.S.O., Nairobi.							
Nerlove, M. (1964),	"Spectral Analysis of Seasonal Adjust- ment Procedures", Econometrica.Vol.32, No.3, p241-286.							
Obasi, G.O.P. and H. Rhode(September,1974),	"Some Factors of the Atmospheric Enviroment in Kenya", Paper presented at the Annual Symposium of the E.A. Academy, Nairobi.							
Blackman, R.B. and J.W. Tukey (1958),	The Measurement of Bower Spectra. Constable and Co., London.							

